

**Environmental Measurements in the Beaufort Sea, Spring 1993** 

by T. Wen and F. Karig



Technical Memorandum
APL-UW TM 4-93
December 1993

This Technical Memorandum has received limited review.





Applied Physics Laboratory University of Washington 1013 NE 40th Street Seattle, Washington 98105-6698

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# Acknowledgments

The research presented in this report was sponsored by the organizations participating in the ICEX 1-93 Applied Physics Laboratory Ice Station (APLIS). Funding was provided by the Naval Sea Systems Command, Code 06UR43. Participating laboratories were the Naval Undersea Warfare Center, New London Laboratory; the Naval Undersea Warfare Center, Arctic Submarine Laboratory; and the Naval Surface Warfare Center, Carderock Division.

The purpose of this report is simply to present the environmental data obtained during the camp. The data analysis is very limited. All the data presented here are stored in digital format and are available for further analysis. Requests for data should be forwarded to

Director Applied Physics Laboratory 1013 N.E. 40th Street Seattle, WA 98105-6698

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## **ABSTRACT**

This report presents environmental data obtained by the Applied Physics Laboratory of the University of Washington (APL-UW) at APLIS 93, an ice camp established in the Beaufort Sea in Spring 1993 to support Navy-sponsored tests and research during ICEX 1-93. Data analysis is limited because the purpose of this report is to provide field data to ice camp participants. The data were collected to document the meteorological and oceanographic conditions that existed during camp activities. The main data sets are weather, floe drift, STD profiles, currents, and ice properties.

#### I. INTRODUCTION

This report presents environmental data taken in the spring of 1993 at APLIS (Applied Physics Laboratory Ice Station) in the Beaufort Sea. The ice camp was established and maintained by personnel from the Applied Physics Laboratory, University of Washington, to support Navy-sponsored research and test activities conducted by the organizations participating in ICEX 1-93. The environmental data — weather, floe drift, STD profiles, currents, and ice properties — were gathered by APL-UW personnel and are intended to support the analysis of experimental data obtained by ICEX 1-93 participants. APL-UW has been conducting acoustic and oceanographic research in the Arctic since the 1970s. The ICEX 1-93 data are therefore part of a growing database for the Beaufort Sea.

APLIS 93 was established on a multiyear floe approximately 275 km north of Prudhoe Bay, Alaska (see Figure 1). The floe was selected on 22 March after a one-day search. Camp setup took place the following week. During this build-up phase, five people erected a mess hall, sleeping quarters, generators, and control building. Collection of environmental data started on 29 March and lasted until 16 April, when the objectives of ICEX 1-93 were met. During the test period, personnel at the camp numbered as many as 35 and dwindled to only a few near the end. Evacuation of the camp was completed on 17 April.

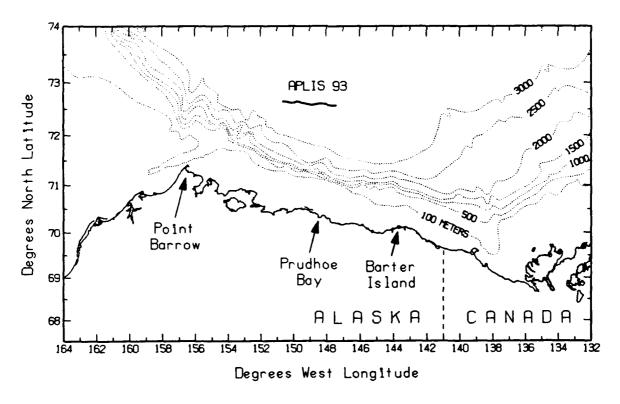


Figure 1. Position of APLIS 93.

For quick access to a runway, a site at the edge of the floe next to a flat refrozen lead was chosen for the camp. The thickness of the ice at the camp and at other level locations of the floe was 2-3 m. The runway lead was about 1 m thick and over 500 m long. An underwater acoustic tracking range consisting of four hydrophones was deployed to track underwater vehicles participating in the ICEX tests.

Air temperature and pressure and wind speed and direction were recorded automatically at 5-minute intervals. The weather was generally calm, except for a 24-hour period during which high winds built up snow drifts in and around the camp.

Because the floe drifted in response to wind stress, its position needed to be tracked for the analysis of some data such as satellite imagery. Latitude and longitude of the floe supplied by a GPS receiver were logged at 30-minute intervals and also used to determine drift speed and direction.

CTD casts were made often to determine the properties of the water column down to 350 m. Sound speed profiles were then derived from the measured temperature and salinity and were used to predict the real-time performance of acoustic equipment and the underwater tracking range.

The time of acquisition for some data is tagged "L" for local time, while others are tagged "UTC" for Universal Coordinated Time. Local time is UTC – 9 hours.

Because the purpose of this report is simply to present environmental data from the camp, analysis is very limited. All data presented here are stored in digital format and are available upon request. The Spring 1993 CTD data, as well as those from previous years, are also available from the National Oceanographic Data Center in Washington, D.C.

#### II. THE FLOE

Selection of an ice floe suitable for a camp was based on several requirements. First, we needed a refrozen lead long enough and thick enough (at least 1 m) to serve as a runway, since transportation to and from the camp depended entirely on aircraft (only one Twin Otter this year). Second, the floe had to be over deep water for the long-range underwater tracking operation. This required that the camp be located north of 72° latitude, where the water is at least 3000 m deep. Third, the initial site had to be far enough east to allow for the westerly drift that would (historically) occur during the period of camp occupancy. As in previous years, the first criterion narrowed down the usable floes to those north of 72°N so the second criterion was fulfilled automatically. This year, a suitable floe was found at 72.5°N and 147.1°W after one day of air search.

The camp was established at the edge of the floe, next to a refrozen lead (see Figure 2). Ice in the floe, excluding hummocks, varied from 2 to 3 m thick, with a snow cover of 5–10 cm. Ice in the refrozen lead was 1 m thick with the same amount of snow cover. A 550-m runway was constructed on the lead by leveling off the uneven snow cover with a snowmobile-towed grader. Figure 3 shows an ERS-1 C-band microwave SAR image of the ice pack. The X marks the location of the camp. In this image, made on 15 April, light areas represent multiyear ice and dark areas represent first-year or young ice. The darkest area, such as the one found in the upper right-hand corner, was most likely open water.



Figure 2. Aerial photo of ice camp.

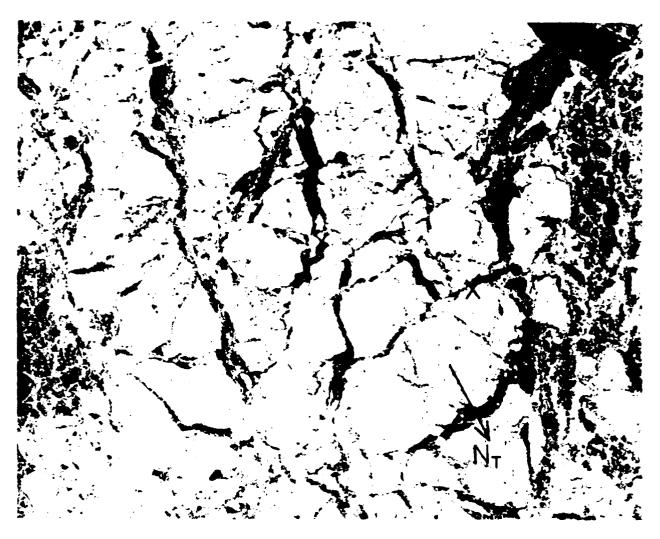


Figure 3. C-band microwave SAR image of 25 km × 25 km area containing the ice floe. ERS-1, Beaufort Sea, 15 April 1993. Dark areas represent mostly first-year ice and some open water. (Copyright European Space Agency; provided by Alaska SAR facility)

The ICEX 1-93 tests involved submarines. To track the submarines, we set up an underwater tracking range with an imaginary X-Y coordinate system shown in Figure 4. The origin of the coordinate system was a 1-m-diameter hydrohole (marked "O" on the figure) beneath the control building. Four hydrophones, each approximately 500 m from the camp, were suspended at 30.5-m depth for receiving cw pulses transmitted by the submarines. The signals were cabled back to the control building, filtered, amplified, validated, and time stamped. An x,y fix can then be made based on the time-of-flights and an average sound speed. The operational distance of the range is 15,000 m, sometimes as far as 20,000 m, but with large scatter in the x,y fixes. To transform an x,y fix in

the arbitrary X-Y system to earth coordinates, the latitude and the longitude of the origin and the true bearing of the +Y axis must be known. The true bearing of the +Y axis was determined by celestial sightings and was found to vary from 84.8° to 82.3° due to floe rotation.

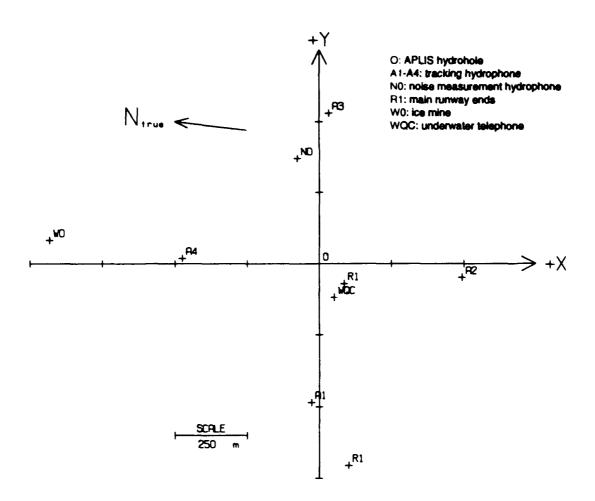


Figure 4. Underwater tracking range X-Y coordinate system.

The magnetic variation at the general locale was 31°, calculated using GEOMAG, a PC program published by the Naval Oceanographic Office and based on the World Magnetic Model for Epoch 1990 (WMM-90).

#### III. FLOE MOVEMENT

The camp's position was determined using a Global Positioning System (GPS) receiver (Kinemetrics/Truetime GPS-DC) and displayed and logged on a PC computer via RS-232 bus. GPS fixes were read from the receiver every 30 minutes and stored on a floppy disk. At least four GPS satellites were in view 99% of the time. Consequently, the geometric dilution of precision, an indicator of tracking precision, was usually 3 or less. But because of the uncertainty of up to 100 m in the fixes, the data had to be smoothed using an 11-point binomial filter during postprocessing. The smoothing operation is essential when floe drift speed and direction are to be determined. Figure 5 shows the drift track of the floe. During the camp period, the floe drifted a net distance of 118 km toward the west. Drift speed and direction are plotted in Figure 6 along with wind speed and direction (sense reversed); the plots show good correlation between the wind and drift. As a rule of thumb, the floe drift speed is about 1/60-1/50 of the wind speed, sometimes even less, and its direction is about 45° to the right of the wind direction. The floe's position, as well as its drift speed and direction, is listed in Appendix A.

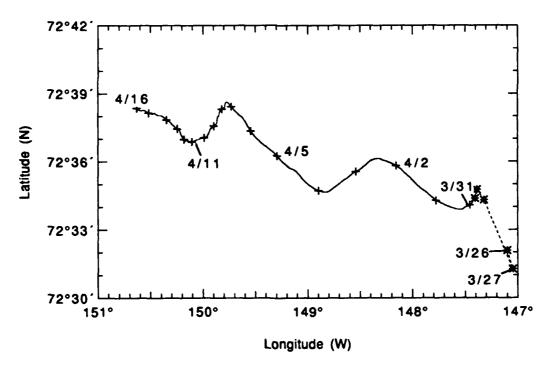


Figure 5. Drift track of APLIS 93. Asterisks denote positions obtained with a portable GPS receiver.

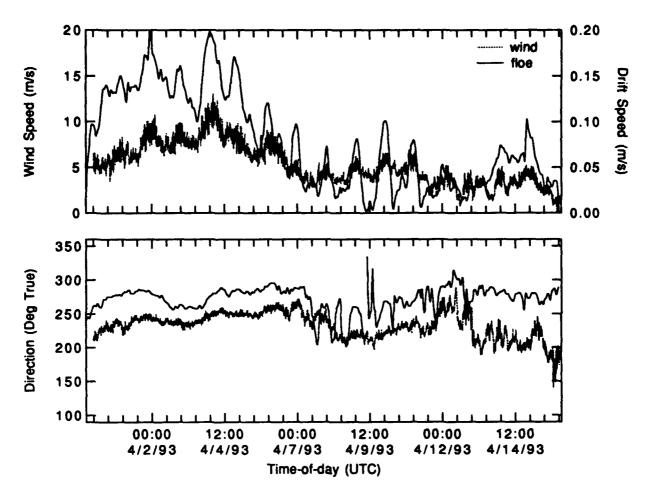


Figure 6. Floe drift speed and direction, with wind data superposed for comparison.

Our GPS receiver also provided 1-Hz timing pulses accurate to within microseconds of UTC. These were used to check the drift of the tracking range master clock and were therefore crucial to the precision of our underwater tracking range.

The rotation of the floe was determined by measuring the change in the true bearing of the +Y axis of the X-Y coordinate system. To obtain the true bearing of the +Y axis, the grid bearing of the Sun was first read with a transit set up on top of the control building over the hydrohole at (0,0). The time at which the bearing was read was noted to the nearest second. Using the time and current location, the true bearing of the Sun was calculated using the PC program MICA (Multiyear Interactive Computer Almanac; U. S. Naval Observatory, Washington, D.C.). The difference between the grid and true bearings of the Sun was the true bearing of the +Y axis. Celestial sightings were made frequently until the last few days when overcast conditions precluded sightings. The +Y axis of the coordinate system varied from 84.8° to 82.3° during the observation period (see Table 1), indicating a CCW floe rotation of 2.5°.

Table 1. Bearing of the +Y axis of underwater tracking range. All sightings were made on the Sun with the exception of one on Venus, flagged with \*.

Time mmdd	(UTC) hhmmss	Lat. dd mm ss	Long. ddd mm ss	Azimuth of Sun	Grid Bearing	True Bearing of +Y Axis
0401	002926	72 34 19	147 47 07	220.71	135.87	84.84
	005330	72 34 21	147 47 34	226.93	142.02	84.91
0403	194645	72 34 39	148 48 48	144.85	60.73	84.12
	195017			145.77	61.67	84.10
0404	231933	72 36 13	149 17 12	201.45	116.62	84.83
	232205			202.12	117.31	84.81
0406	000120	72 37 21	149 32 42	212.39	127.52	84.87
	000707			213.91	129.02	84.89
0407	021226	72 38 32	149 45 02	245.74	160.92	84.82
	021448			246.32	161.52	84.80
0408	175538	72 37 45	149 52 02	115.46	30.83	84.63
	175915			116.36	31.71	84.65
0409	201601	72 37 11	149 57 25	151.44	67.79	83.65
	201925			152.34	68.70	83.64
	235259	72 37 04	149 59 22	210.26	126.84	83.42
	235643			211.25	127.83	83.42
0410	074946	72 36 57	150 02 27	156.86	73.64	83.22*
	173715	72 36 58	150 04 15	110.65	27.75	82.90
	174016			111.39	28.53	82.86
	214558	72 36 54	150 05 26	175.75	93.07	82.68
	214910			176.63	93.97	82.66
0411	235358	72 36 59	150 10 59	210.58	128.22	82.36
	235711			211.44	129.08	82.36

## IV. WEATHER

Weather information was important for research as well as for logistics. There was also general interest in the wind speed and air temperature. Weather parameters therefore became one of the routine environmental measurements made at APLIS 93. Data were collected by a weather station (Weatherpak 100/Coastal Climate Co.) mounted on a telescoping mast at a height of 10 m. The mast was guy-wired so that it would not sway in high winds and cause erroneous wind speed readings. Air temperature, atmospheric pressure, and wind speed and direction were measured and recorded at regular intervals. The accuracy of the meteorological measurements is 0.5 m/s for wind speed, 2° for wind direction, 0.5 mbar for atmospheric pressure, and 0.2°C for temperature.

A laptop PC in the control building was used to control the weather station and to display meteorological data. Commands and data were sent over an RS-232C link. The Weatherpak was programmed to average the weather parameters for 5 s at 5-minute intervals. Sampled data were stored in the Weatherpak's internal memory and also downloaded to the PC for display and storage on floppy disk. Weather parameters could also be read at any time by pressing a programmed menu key on the PC.

The data are plotted and presented in Figure 7. Wind direction was originally logged as magnetic but has been converted to true using an average variation of  $31^{\circ}$ . The temperature and the wind speed showed diurnal variations until the overcast condition set in. With the overcast, the temperature hovered about  $-10^{\circ}$  C. During this period of high temperatures and low winds, the weather seemed very warm, compared to earlier times.

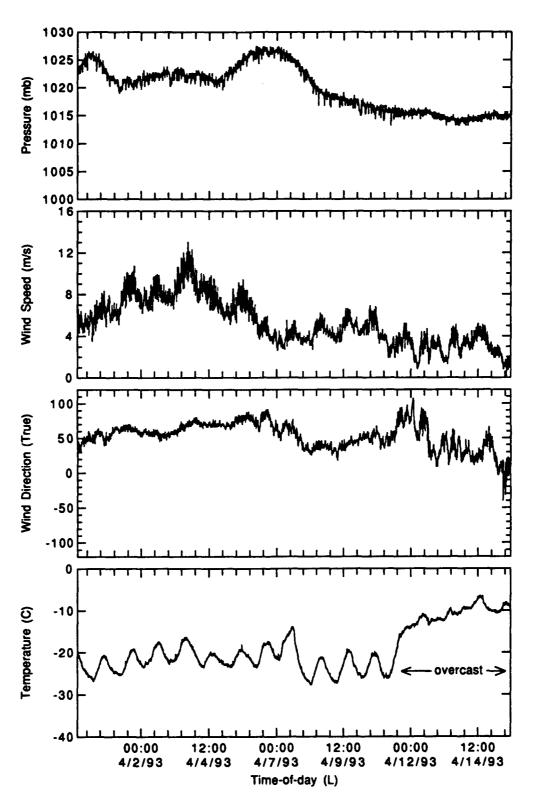


Figure 7. Weather parameters.

#### V. CTD MEASUREMENTS

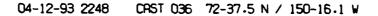
CTD casts were made daily to obtain the temperature and salinity properties of the water column. The components of the CTD profiler were a solid-state data logger (Sea-Bird), a thermistor (Sea-Bird), a conductivity cell (Sea-Bird), and a pressure sensor (Paroscientific Digiquartz). The profiler was attached to the end of a 6.4-mm-diameter nylon line and deployed with an ac-powered winch. To ensure adequate flushing of seawater through the conductivity cell, the profiler was allowed to free-fall (attached to the line), reaching a speed of ~1.3 m/s. With the sampling rate of the logger at 8 Hz, the water column was consequently sampled at ~16 cm intervals, resulting in high-resolution temperature and salinity profiles. The casts were generally to only ~350 m or less because water properties at greater depths do not vary significantly from day to day. After each cast, the raw data were read out of the logger via an RS-232C link to a laptop PC for processing and plotting. The raw data were first converted to temperature, conductivity, and depth using sensor calibration constants. UNESCO '83 algorithms<sup>4</sup> were then used to compute salinity, sound speed, and  $\sigma_i$  (the density of the *in situ* water with the pressure reduced to atmospheric). The conductivity cells and depth sensors were last calibrated 24 months before the field trip, and the thermistors 16 months before. Since these sensors are very stable, according to the historical calibration data, we felt the last set of calibration constants was still fairly good. For the measured properties, we suggest an accuracy of 0.005°C for temperature, 0.008 mS/cm for conductivity, and 0.05 m for depth; the computed properties should have an accuracy of 0.01 ppt for salinity, 0.02 m/s for sound speed, and 0.01 kg/m<sup>3</sup> for  $\sigma_t$ .

Table 2 lists the casts made at APLIS 93. The STD plots from all casts are given in Appendix B. Figure 8 shows the STD profiles from a typical cast. The well-mixed upper layer in this case was 40 m deep, although it has extended as deep as 60 m in other years. An intrusion of warmer water from the Bering Sea lies under the mixed layer, creating a thermocline and a halocline (and therefore a pycnocline). This intrusion typically extends to a depth of 100 m. Below this is the colder Chukchi Sea water. From -200 m down is Atlantic Water with a temperature maximum of 0.5°C. A temperature staircase is observed below 250 m with an average step change of 1-2 m and 0.01-0.02°C. The temperature and salinity profiles are presented in Figures 9a and 9b, respectively, in waterfall format to show the temporal as well as spatial variations as the floe drifted about.

The thermocline/halocline at 30-40 m depth could have had a large effect on the underwater tracking. Figure 10 shows a ray trace based on the sound speed profile of cast #22. It can be seen that there is a dearth of rays in the weak zone bounded by 2000-4500 m in range and 0-100 m in depth. A source in this zone would have been more difficult to hear through the hydrophones compared with a source outside the zone. We experienced this phenomenon when we tried and had great difficulty in obtaining the x,y fix of some hydroholes at these ranges using a pinger at various depths to 30.5 m. Tracking of submarines was generally trouble-free, as they usually stayed below 100 m depth and out of the weak zone.

Table 2. CTD casts made at APLIS 93.

Date	Time (Local)	Cast No.	Latitude	Longitude
02 20 02	1051	C + Cm 00 +	7000400	
03-29-93	1851	CAST 001	72°34.8′N	147°23.3′W
03-30-93	0649	CAST 002	72°34.5′N	147°24.2′W
03-30-93	1738	CAST 003	72°34.0′N	147°29.5′W
03-31-93	0632	CAST 004	72°34.0′N	147°39.1′W
03-31-93	2016	CAST 005	72°34.6′N	147°51.3′W
04-01-93	1613	CAST 006	72°35.9′N	148°11.2′W
04-02-93	0630	CAST 007	73°36.0′N	148°24.8′W
04-02-93	1520	CAST 008	72°35.6′N	148°32.5′W
04-02-93	1847	CAST 009	72°35.4′N	148°36.1′W
04-02-93	2229	CAST 010	72°35.2′N	148°38.9′W
04-03-93	0630	CAST 011	72°34.9′N	148°44.6′W
04-03-93	1406	CAST 012	72°34.7′N	148°52.5′W
04-03-93	2128	CAST 013	72°35.1′N	149°01.3′W
04-04-93	0610	CAST 014	72°35.6′N	149°08.5 ′W
04-04-93	1416	CAST 015	72°36.1′N	149°16.2′W
04-05-93	0848	CAST 016	72°36.7′N	149°28.5′W
04-05-93	1654	<b>CAST 017</b>	72°37.3′N	149°32.7′W
04-06-93	0934	<b>CAST 018</b>	72°38.1′N	149°40.7 °W
04-06-93	1147	CAST 019	72°38.2′N	149°41.5′W
04-06-93	1617	<b>CAST 020</b>	72°38.5′N	149°44.2′W
04-06-93	2140	CAST 021	72°38.6′N	149°45.8′W
04-07-93	0751	<b>CAST 022</b>	72°38.5′N	149°47.5′W
04-07-93	2014	<b>CAST 023</b>	72°38.1 'N	149°50.7′W
04-08-93	0914	CAST 024	72°37.8′N	149°51.9′W
04-08-93	1844	CAST 025	72°37.5′N	149°55.4′W
04-09-93	0831	CAST 026	72°37.3′N	149°56.5′W
04-09-93	1427	CAST 027	72°37.3′N	149°58.3′W
04-09-93	2239	CAST 028	72°36.9′N	150°02.5′W
04-10-93	1232	CAST 029	72°36.9′N	150°04.9′W
04-10-93	1750	CAST 030	72°36.9′N	150°07.3′W
04-11-93	0656	CAST 031	72°36.9′N	150°09.5 'W
04-11-93	1424	CAST 032	72°36.9′N	150°10.5 'W
04-11-93	2340	CAST 032	72°37.2′N	150°10.5°W
04-12-93	0746	CAST 033	72°37.4′N	150°13.8′W
04-12-93	1450	CAST 034	72°37.4′N	150°13.6°W
04-12-93	2248	CAST 035	72°37.5′N	150°14.0°W
04-12-93		CAST 030	72°37.3′N	150°18.0°W
04-13-93	0754		72°38.0′N	
	2320	CAST 038		150°24.0′W
04-14-93	0635	CAST 039	72°38.1′N	150°26.7′W
04-14-93	1418	CAST 040	72°38.1′N	150°30.2′W
04-14-93	2134	CAST 041	72°38.3′N	150°33.7′W
04-15-93	0621	CAST 042	72°38.8′N	150°36.1 ′W
04-15-93	1407	CAST 043	72°38.3′N	150°37.7′W



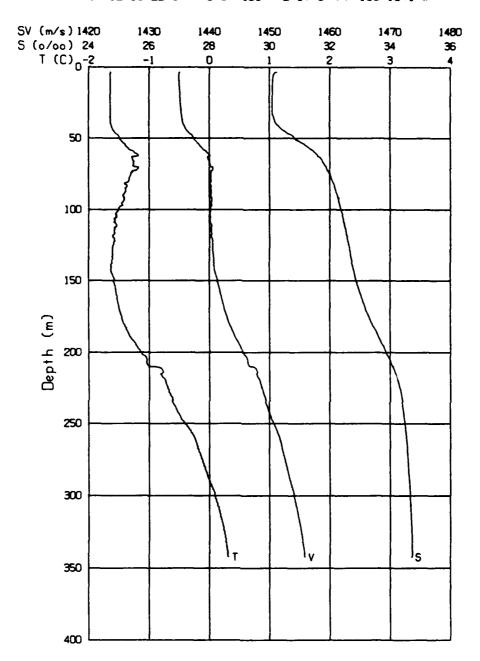


Figure 8. Sample STD profiles from cast 36 showing intrusions at 70 m and 220 m depth and temperature staircase below 250 m.

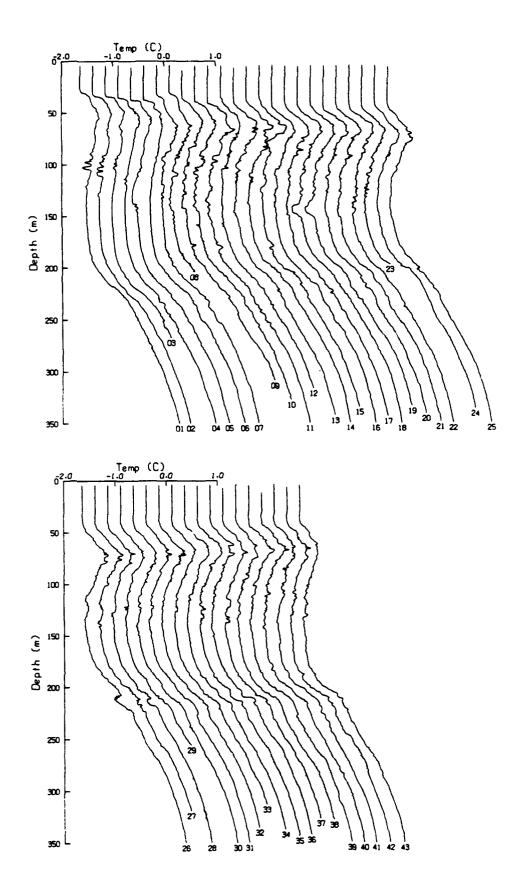


Figure 9a. Waterfall presentation of temperature profiles.

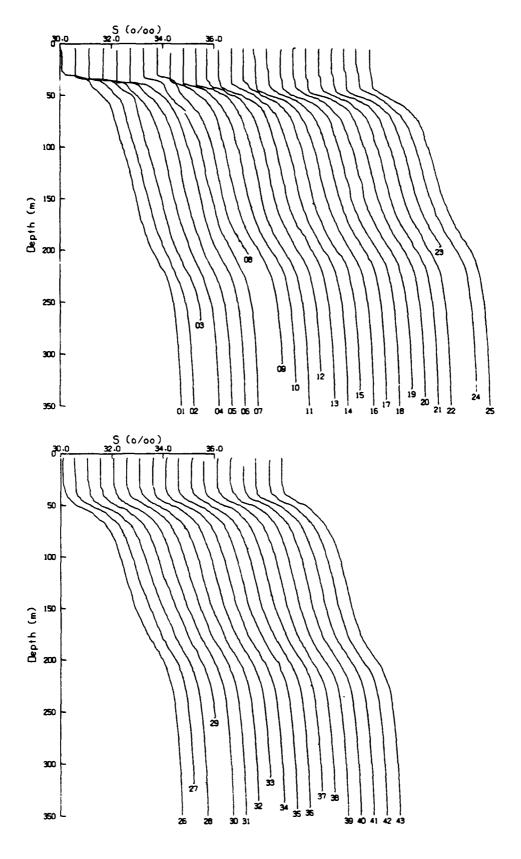


Figure 9b. Waterfall presentation of salinity profiles.

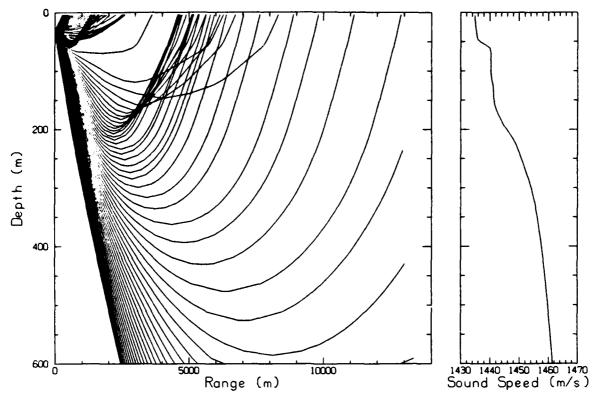


Figure 10. Ray trace showing weak zone, cast #22. Source depth at 30.5 m, rays from -15° to 5° at 0.2° increment.

#### VI. CURRENTS

To measure current, the CTD profiler was replaced with an InterOcean S4 current meter. A 4.5-kg lead weight was hung 1 m below the instrument to reduce streaming. The default sampling rate of the current meter was 2 Hz, but generally we recorded 5-s (10 samples) or 60-s (120 samples) averages. In addition to recording the magnetic north and east components of the current, the S4 also recorded depth. To obtain stable current measurements, we lowered the meter at depth increments of about 5–10 m, stopping for approximately 1 minute or longer at each depth. When the current meter was brought back to the surface, the data were downloaded to a computer for processing. The mean and standard deviations were computed for the data at each depth. Time series at selected depths, based on CTD information, were also made to assess the temporal variation. Table 3 lists the current meter casts. Vertical current profiles and time series data are shown in Appendix C. Both relative and absolute currents are shown for the vertical profiles, with the former plotted as a solid line and the latter dashed. Bars plotted at each depth represent plus or minus one standard deviation about the averages. In general, the currents were weak. A shear layer is observed at about 205 m in casts 3 and 4.

Table 3. Current meter casts made at APLIS 93.

mmdd	hhmm (Local)	Cast	Floe Drift Ave (s)	cm/s	Dir	Comments
0401	1635	1	60	17.3	278	
0404	1810	2	60	12.0	281	Time series at 70 m
0405	1745	3	5	10.0	295	
0408	1145	4	5	4.8	254	
0409	1645	5	5	9.9	268	
0410	1815	6	5	4.9	280	
0411	1456	7	5	5.3	289	Time series at 225 m
0413	0915	8	5	5.0	286	Time series at 220 m

#### VII. ICE CORE SAMPLES

Some tests were conducted in first-year ice several kilometers from the camp. The ice ranged from 0.6-1 m in thickness. Cores were taken from this ice to determine the salinity and ultimately its physical properties. Some samples were also removed for analysis from refrozen melt ponds and an ice mine.

A 7.6-cm (inside diameter) SIPRE corer was used to remove short segments about 0.5 m in length. Each segment was placed in a miter box and cut into sections 7.5 cm long. Each section was then sealed in a Ziploc bag and tagged. This procedure was repeated until the whole segment, and ultimately the whole ice column, was sampled. The temperature of each section was not measured due to difficulties in handling and the unavoidable cooling of the ice when exposed to the air during handling. Instead, a linear temperature profile in the ice was assumed by using a line defined by the surface and water temperatures. This assumption results in errors of up to 1°C and seemed reasonable based on past measurements of first-year ice. We emphasize that the temperature profile in the ice is not linear throughout the year but is close to being so at this time of year.

The samples were transported back to camp and allowed to melt at room temperature. We set up a "salinometer" with the thermistor and conductivity cell of the CTD profiler; this configuration allowed us to use the standard data-acquisition and reduction system. The sensors were mounted on a frame. The cell was tilted to avoid trapping air bubbles. A length of Tygon tubing was attached to each end of the glass tubing, which was 1 cm in diameter and 18 cm in length. The thermistor was inserted in the Tygon tubing near the conductivity cell. When the melted sample was shaken and poured into the two types of tubing and the cell, conductivity and temperature were measured simultaneously, and salinity was computed. Prior to analysis, the samples and the sensors were

placed on a bench top for several hours to bring them to the same temperature. This was necessary because any large temperature differential would have affected the conductivity results. To reduce the temperature differential further, the water sample was mixed within the cell by raising and lowering the pinched-off Tygon tubing at one end to pass the water sample back and forth.

Once the salinity of the ice was known, its other physical properties were derived using empirical models. Density, as well as brine volume, was calculated from measured salinity and assumed temperature using relationships derived by Cox and Weeks<sup>5</sup>. We chose not to obtain the density by weight and volume because of the problems of irregular core diameters and brine drainage. This choice also simplified the whole operation considerably by eliminating the need for weighing and measuring the core sections. Young's modulus was computed using the empirical formula of Langleben and Pounder<sup>6</sup>.

#### A. First-Year Ice

Cores were taken from four test sites in the first-year ice and designated by the date of test as shown in Table 4. Plots for the properties of the cores from the first-year ice are given in Appendix D. Salinity profiles from days 4-11 are superposed and shown in Figure 11a. The profiles are typical for first-year ice: the salinity is high near the surface, owing to a faster freezing rate, and decreases as the ice grows thicker and more slowly, allowing a longer time for brine expulsion. The salinity should increase again near the bottom because the brine has not had enough time to drain out, and brine has a higher salinity (for example, 37.6 ppt at  $-2^{\circ}$ C and 70.6 ppt at  $-4^{\circ}$ C) than that of seawater (-30 ppt). However, the measurements show relatively low salinities near the bottom. This could be explained by brine drainage occurring during core retrieval due to seawater dilution or displacement of brine.

Table 4. Summary of ice cores.

Core No.	Thickness (cm)	Surface Temp. (°C)	Comments
4-2 #1, #2	102	-13.0	1.6 m apart
4-5 #1, #2	74	-12.7	> 2 m apart
4-6 #1, #2	76	-12.7	> 5 m apart
4-11 #1	84	-13.1	•
#2, #3	85	-13.6	0.5 m apart, >10 m from #1
4-12 #1, #2			Refrozen melt ponds

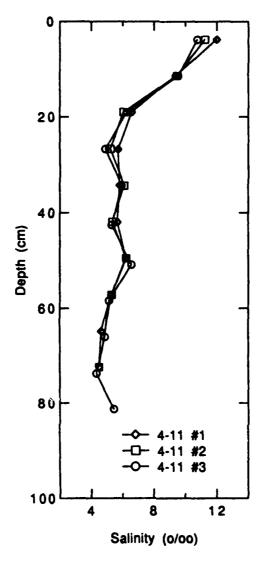


Figure 11a. Salinity profiles in first-year ice, 4-11 #1, #2, and #3 showing interprofile consistency.

Note that in the mid-ice the profile is not smooth, with the salinity varying by as much as 1 ppt from that of the ice above or below. This variation is real and not due to measurement errors, as shown by the consistency of the two cores, #2 and #3, which were taken from ice 0.5 m apart. The vertical variation can be attributed to the different freezing rates that resulted in different lengths of time available for brine drainage.

Figure 11b shows two profiles for cores 4-6 #1 and #2 from the same first-year ice with inter-profile difference of -1 ppt. The cores were taken from locations over 5 m apart. Again, we believe the difference is real and not due to measurement errors, because all sampling procedures were identical and we were able to obtain nearly identical results with cores 4-11 #2 and #3. The horizontal variation in salinity is therefore most likely due to the presence of brine channels and pockets in the ice.

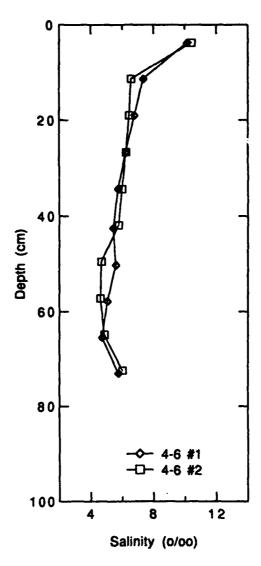


Figure 11b. Salinity profiles in first-year ice, 4-6 #1 and #2, showing interprofile variation.

## B. Melt Ponds and Ice Mine

Figure 12 shows the salinity profiles from two different refrozen meltponds. The salinity is very low in the upper layer. This is expected because when meltwater from ice of 4-8 ppt freezes, an even smaller amount of salt will be trapped in the new ice. Somewhere below 20 cm the salinity jumps up to 4 ppt, comparable to that in typical ice. This tends to suggest that the depths of the melt ponds were at one time on the order of 30 cm.

We also measured the salinity of ice taken from the ice mine that was our source for drinking water. The mine was a 6-m high, well-weathered ridge at least a year old. Two samples were taken from the top of the ridge. The salinity was 0.015 ppt, almost at the noise level of our salinometer. This measurement verifies our judgement that meltwater from multiyear ridge ice tastes as good as fresh water.

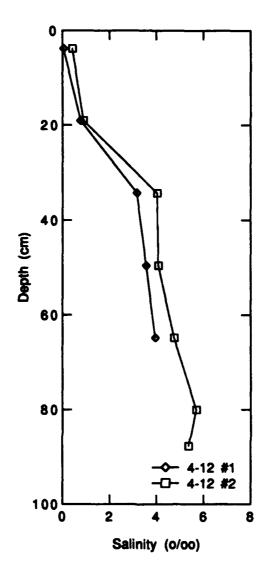


Figure 12. Salinity profiles in two different refrozen melt ponds, 4-12 #1 and #2.

## VIII. REFERENCES

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APPENDIX A
Floe Position and Drift Data
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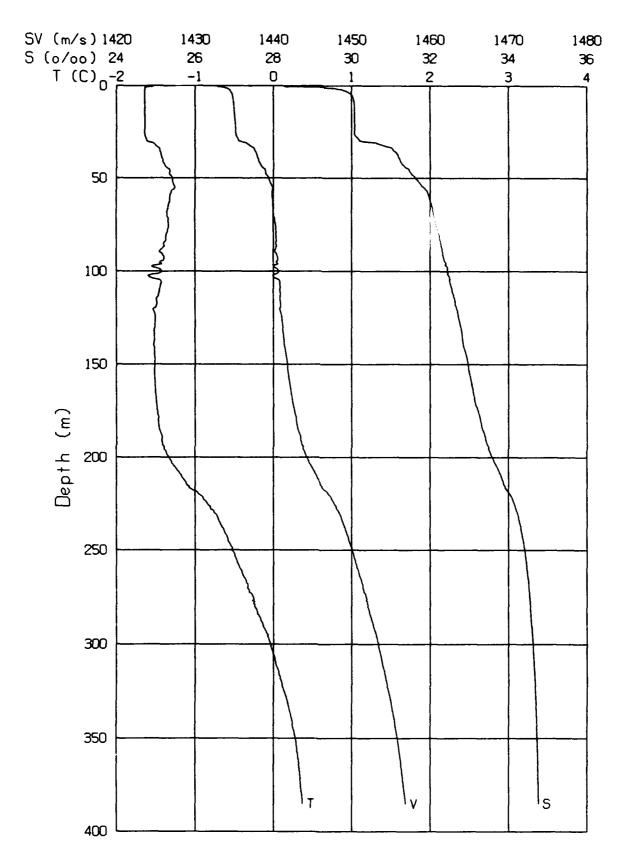
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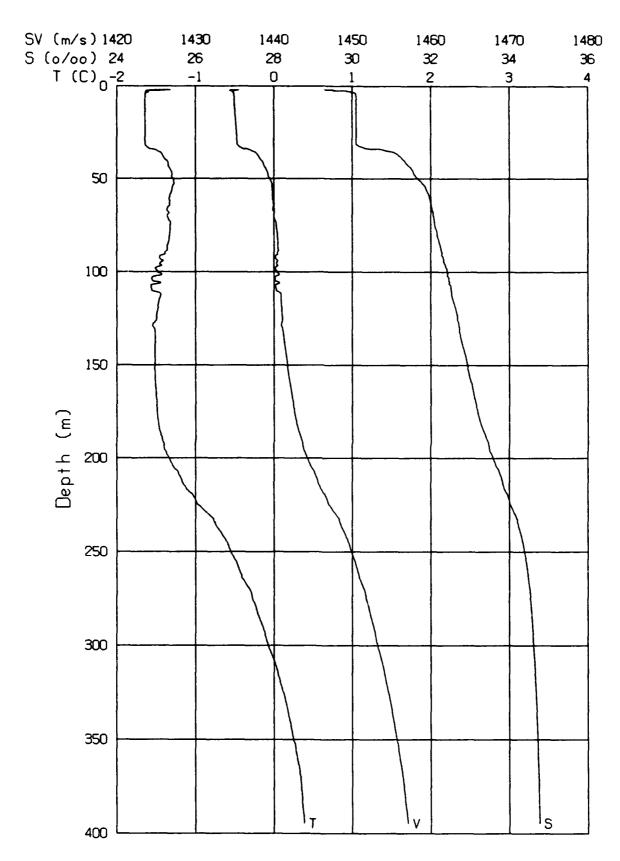
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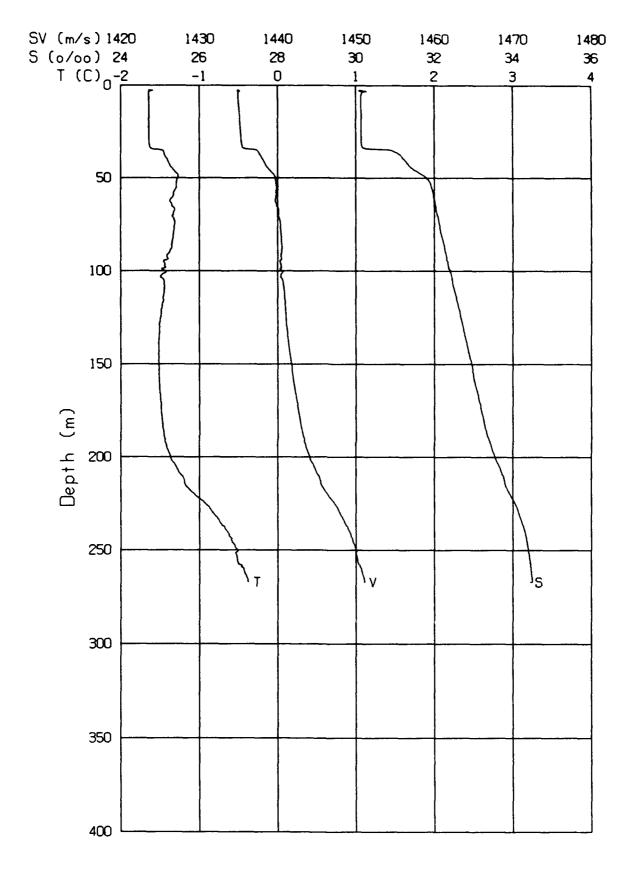
APPENDIX B
STD Plots

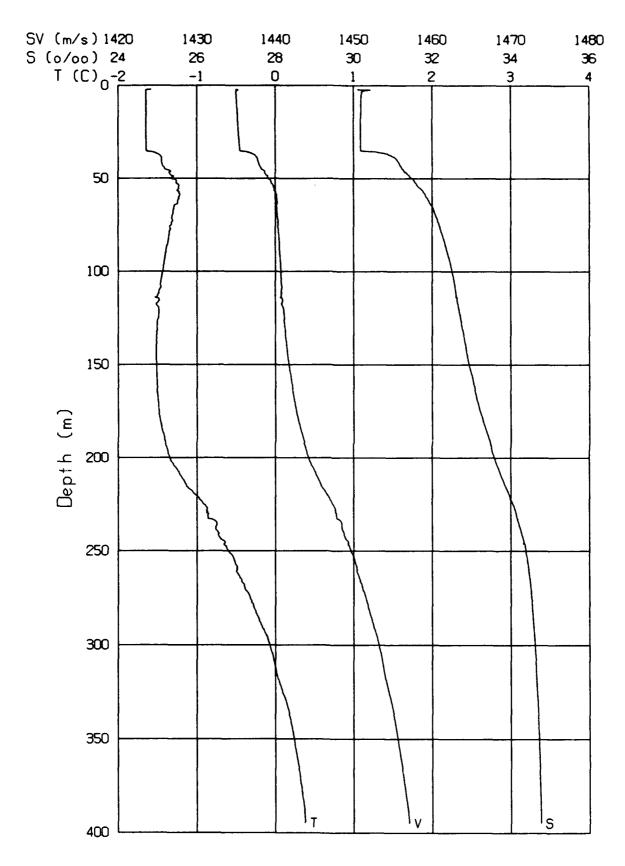
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04-01-93	1613	CAST 006	72°35.9′N	148°11.2′V
04-02-93	0630	CAST 007	73°36.0′N	148°24.8′V
04-02-93	1520	CAST 008	72°35.6′N	148°32.5′V
04-02-93	1847	CAST 009	72°35.4′N	148°36.1′V
04-02-93	2229	CAST 010	72°35.2′N	148°38.9′V
04-03-93	0630	CAST 011	72°34.9′N	148°44.6′V
04-03-93	1406	CAST 012	72°34.7′N	148°52.5′V
04-03-93	2128	CAST 013	72°35.1′N	149°01.3′V
04-04-93	0610	CAST 014	72°35.6′N	149°08.5′V
04-04-93	1416	CAST 015	72°36.1 N	149°16.2′V
04-05-93	0848	CAST 016	72°36.7′N	149°28.5′V
04-05-93	1654	CAST 017	72°37.3′N	149°32.7′V
04-06-93	0934	CAST 018	72°38.1′N	149°40.7′V
04-06-93	1147	CAST 019	72°38.2′N	149°41.5′V
04-06-93	1617	CAST 020	72°38.5′N	149°44.2′\
04-06-93	2140	<b>CAST 021</b>	72°38.6′N	149°45.8′\
04-07-93	0751	CAST 022	72°38.5′N	149°47.5′\
04-07-93	2014	CAST 023	72°38.1′N	149°50.7′\
04-08-93	0914	CAST 024	72°37.8′N	149°51.9′\
04-08-93	1844	CAST 025	72°37.5′N	149°55.4′\
04-09-93	0831	CAST 026	72°37.3′N	149°56.5′\
04-09-93	1427	CAST 027	72°37.3′N	149°58.3′\
04-09-93	2239	CAST 028	72°36.9′N	150°02.5′\
04-10-93	1232	CAST 029	72°36.9´N	150°04.9 1
04-10-93	1750	CAST 030	72°36.9′N	150°07.3′\
04-11-93	0656	CAST 031	72°36.9′N	150°09.5′\
04-11-93	1424	CAST 032	72°36.9´N	150°10.5′\ 150°12.8′\
04-11-93	2340	CAST 033	72°37.2′N	150°12.6 \
04-12-93	0746	CAST 034	72°37.4′N	150°13.6 ′150°14.6 ′1
04-12-93	1450	CAST 035	72°37.4′N	- <del>-</del>
04-12-93	2248	CAST 036	72°37.5′N	150°16.1′
04-13-93	0754	CAST 037	72°37.7′N	150°18.0′
04-13-93	2320	CAST 038	72°38.0′N	150°24.0′
04-14-93	0635	CAST 039	72°38.1 'N	150°26.7′
04-14-93	1418	CAST 040	72°38.1′N	150°30.2′′ 150°33.7′′
04-14-93	2134	CAST 041	72°38.3′N 72°38.8′N	150°33.7
04-15-93	0621	CAST 042	72°38.8 N 72°38.3′N	150°36.1
04-15-93	1407	CAST 043	12°38.3 N	190-37.7

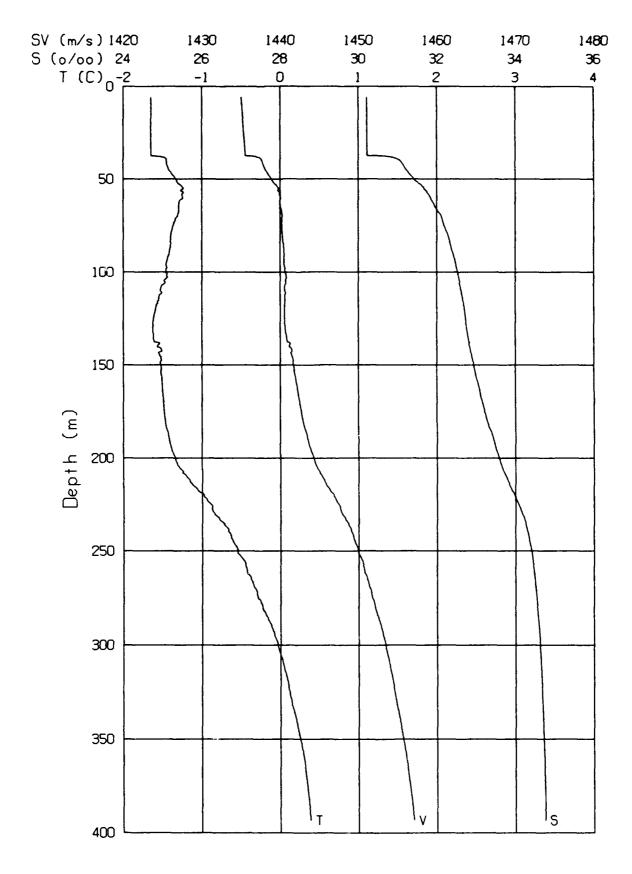


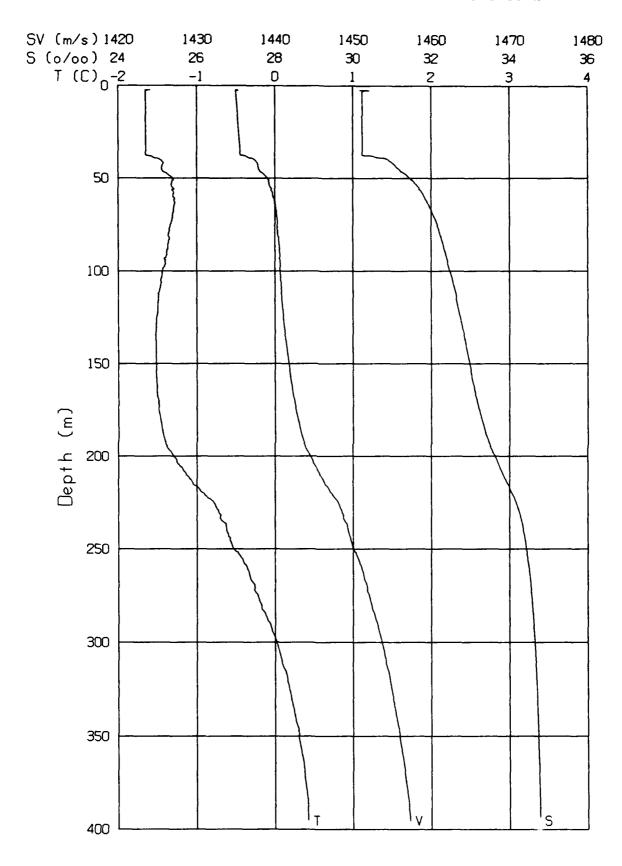


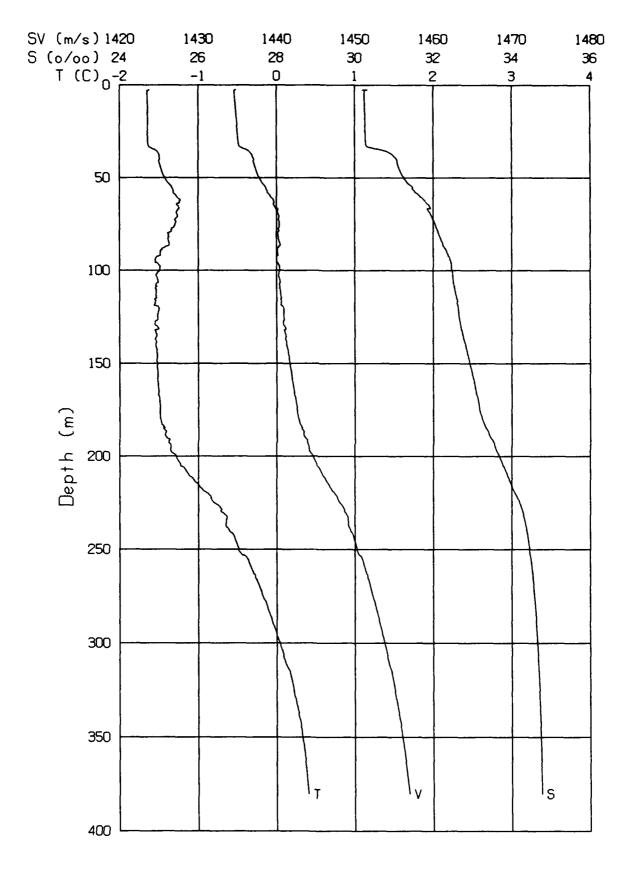
03-30-93 1738 CAST 003 72-34.0 N / 147-29.5 W

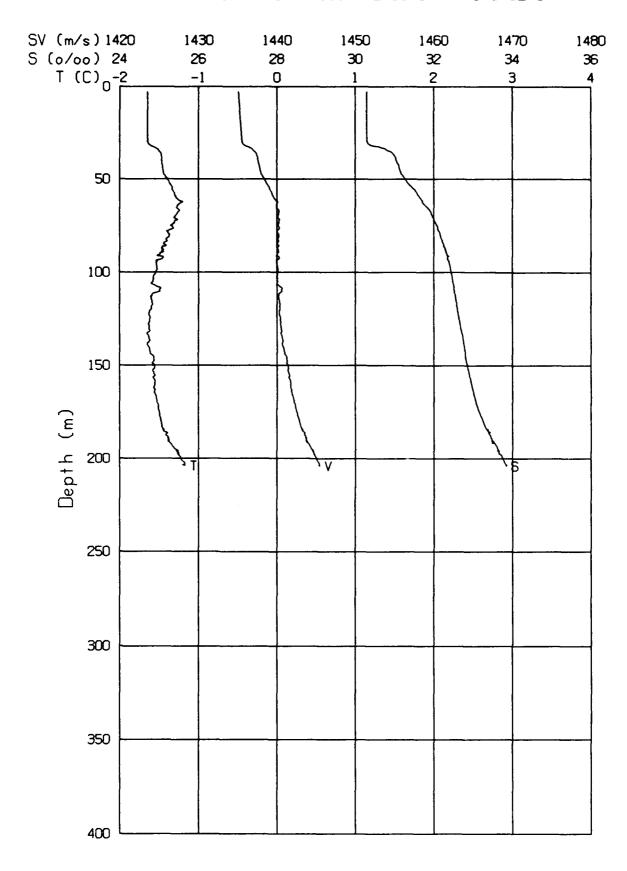


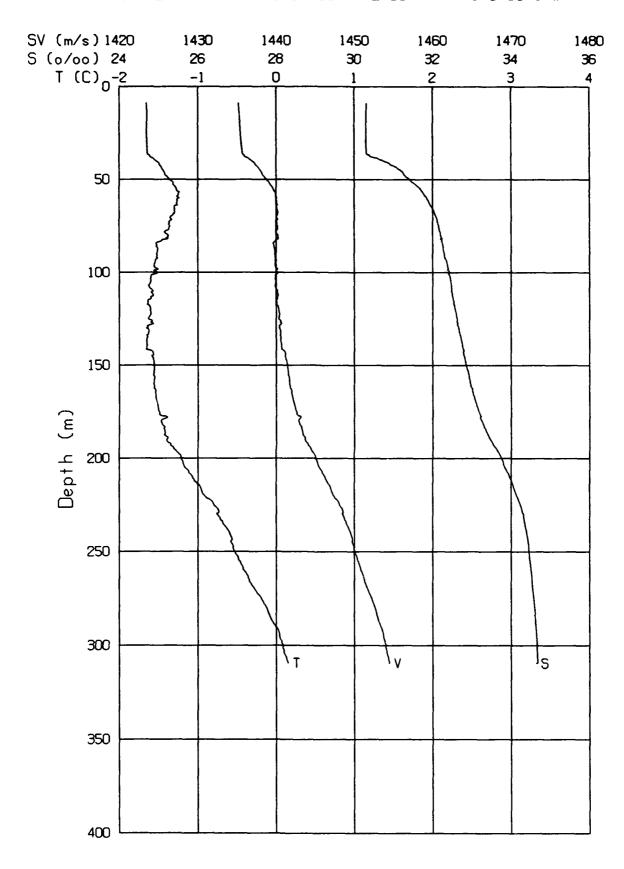


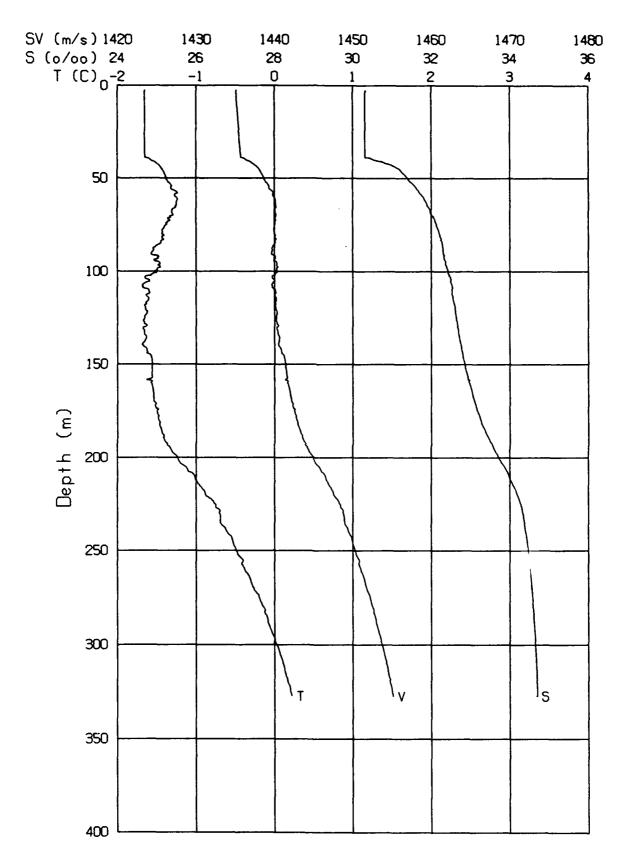


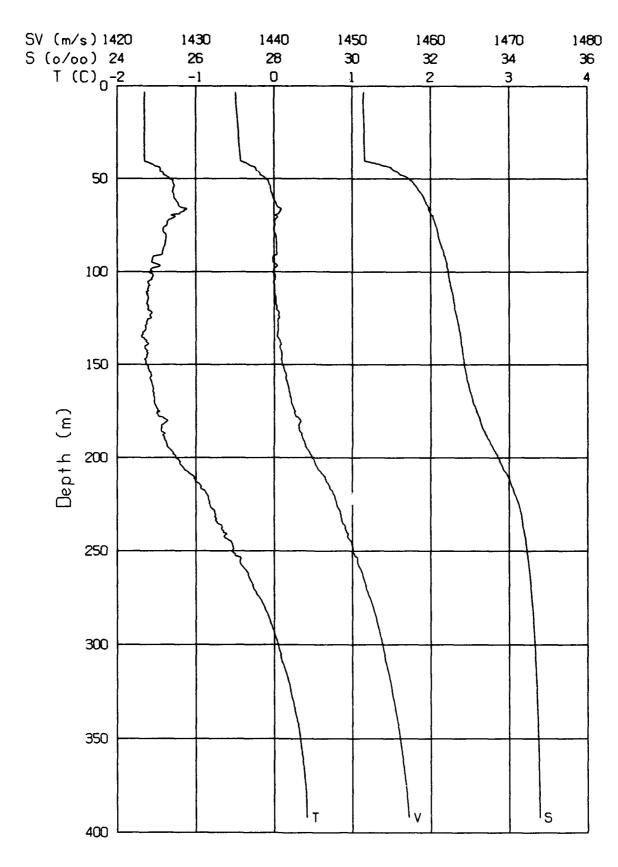


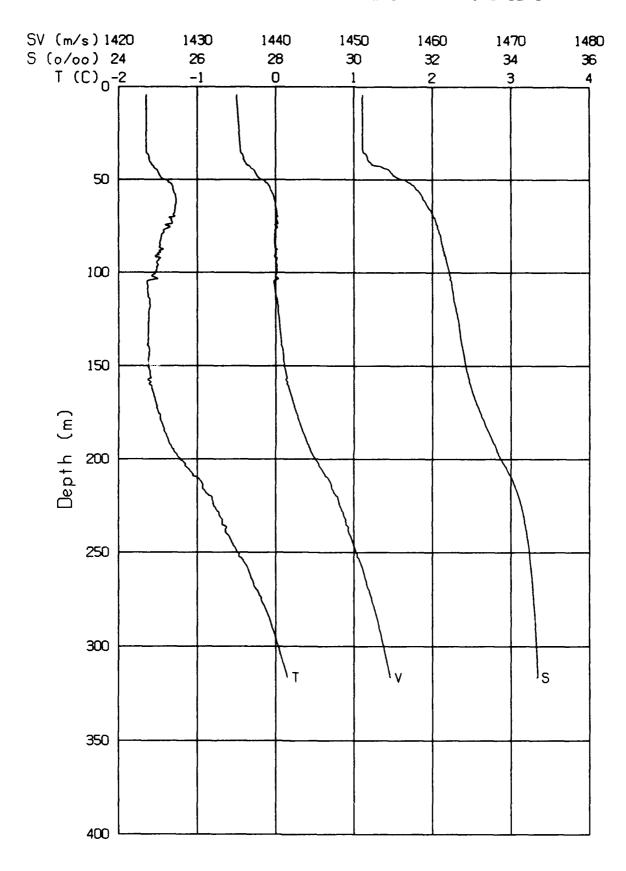


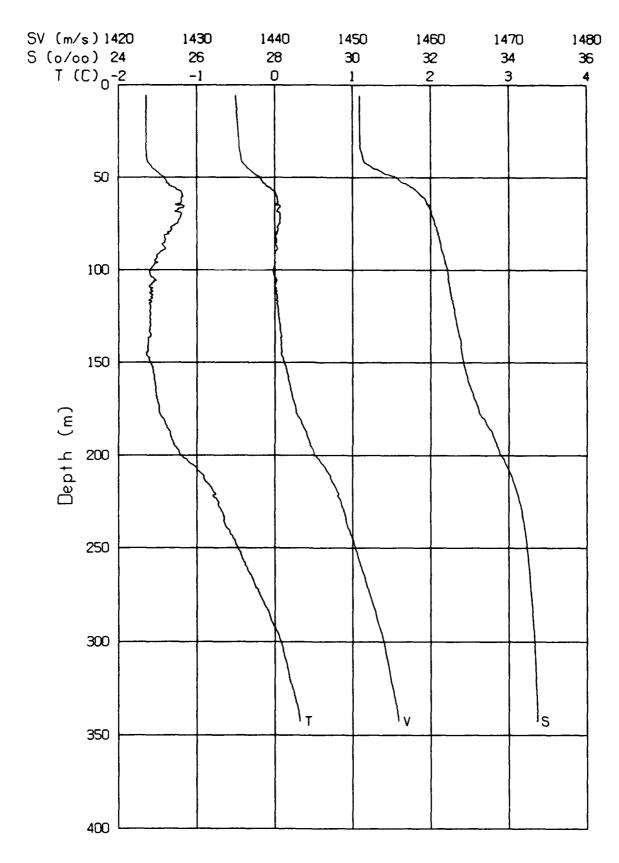


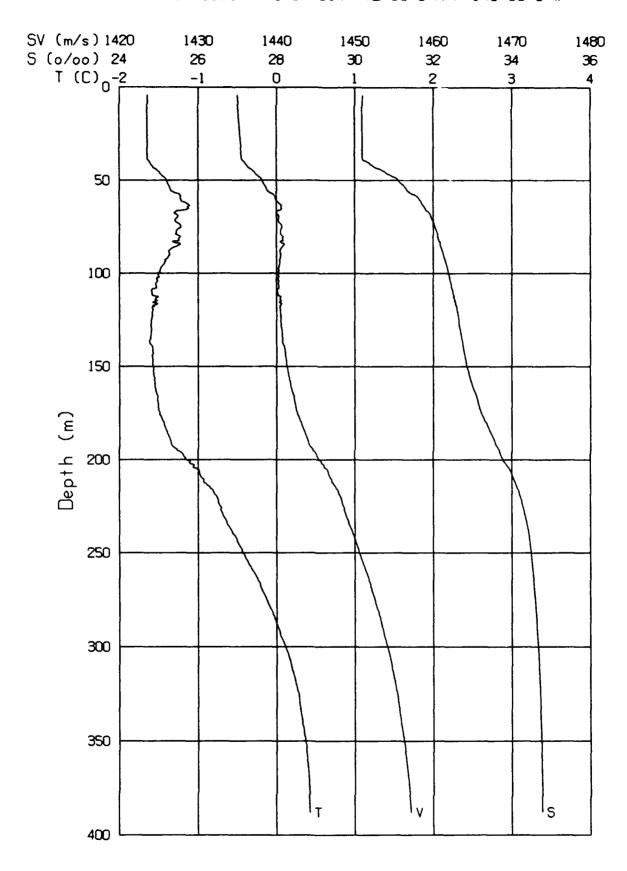


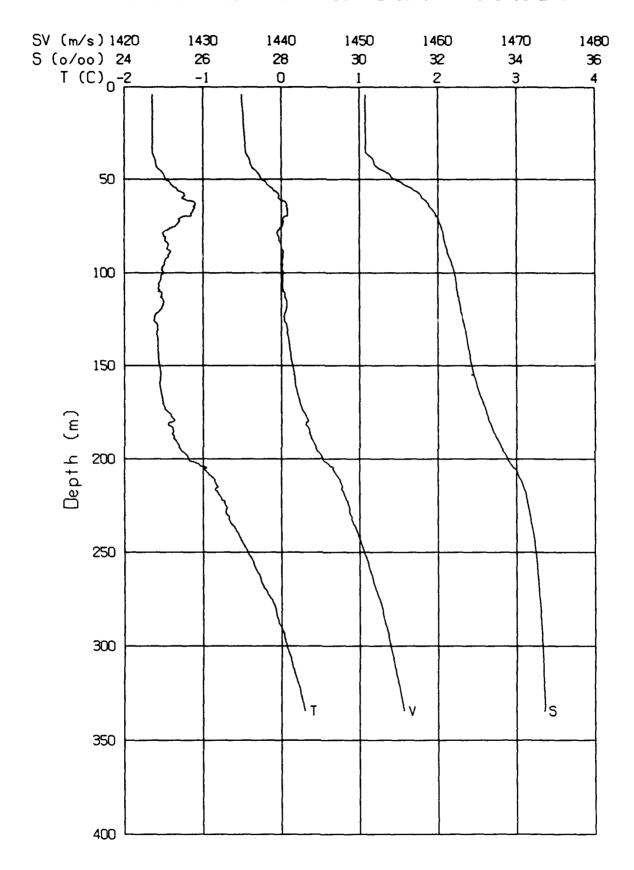


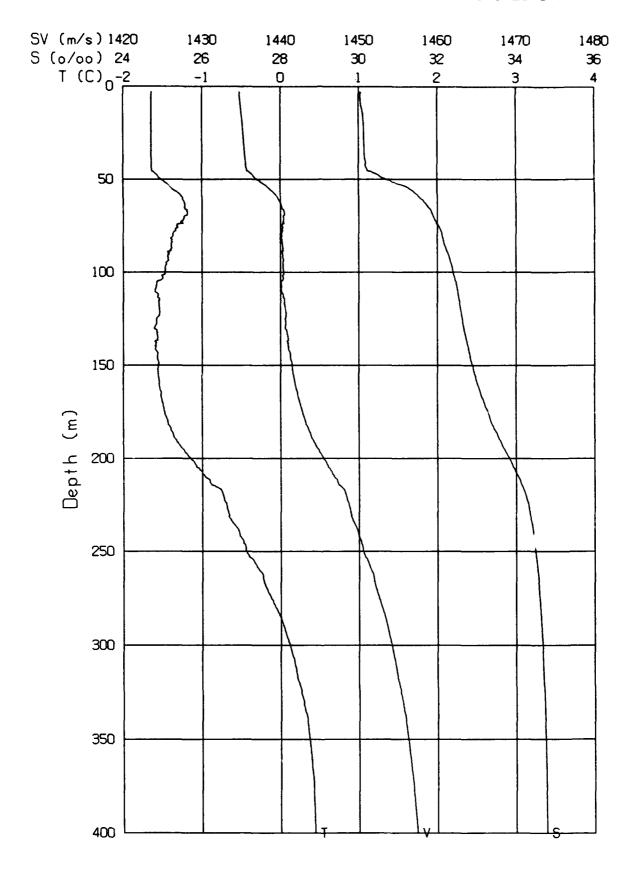


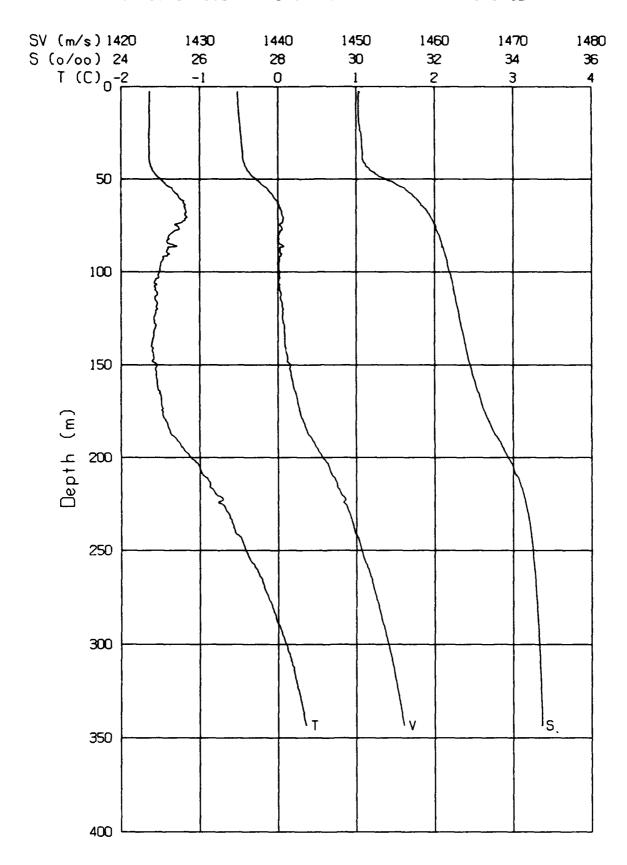


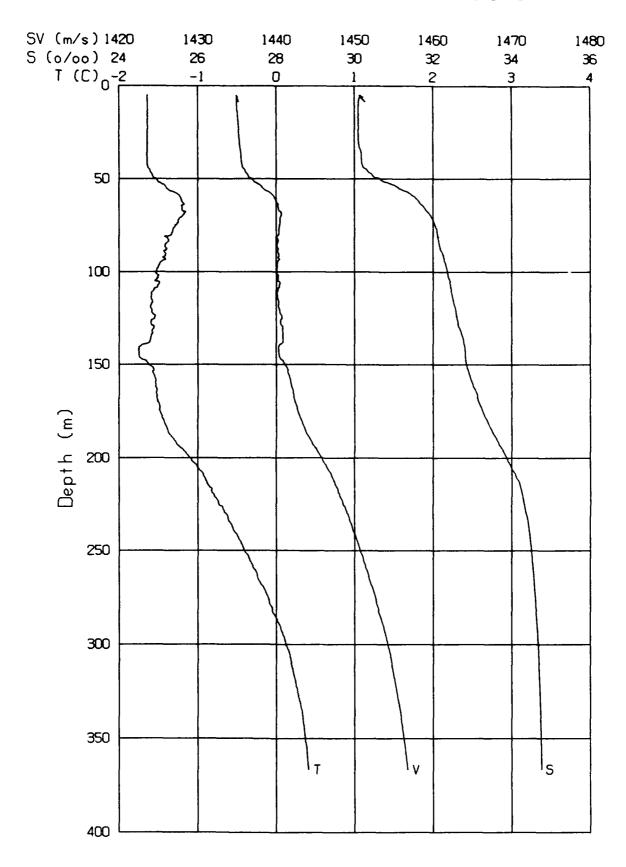


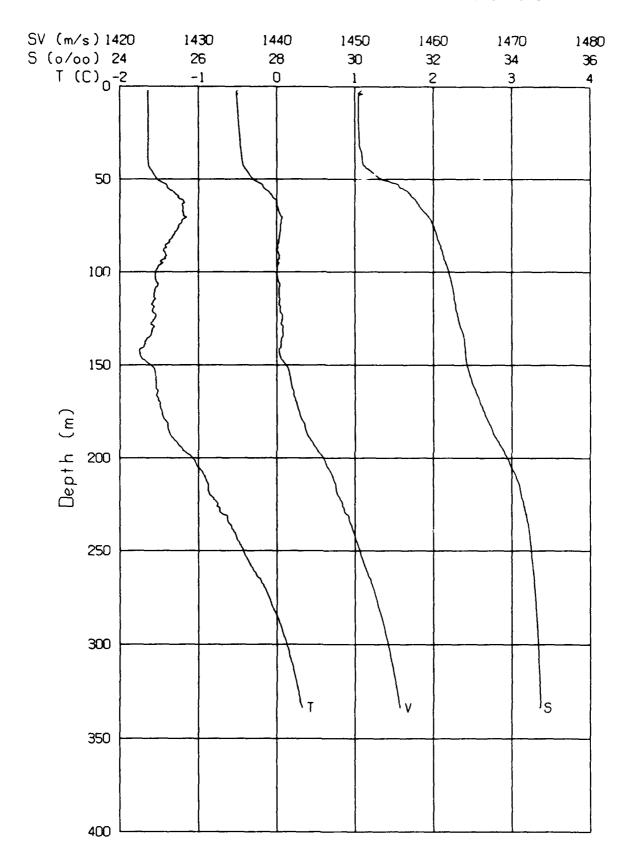


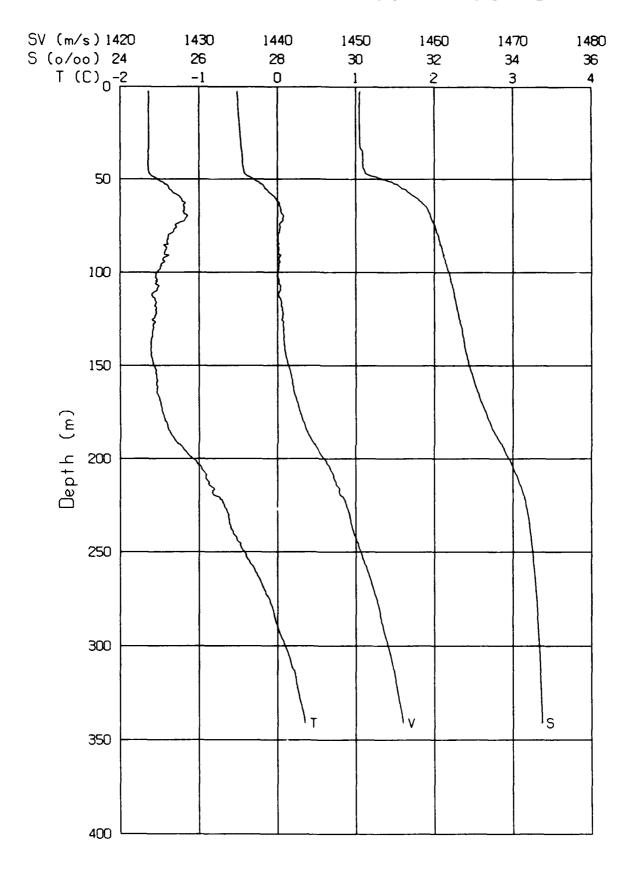


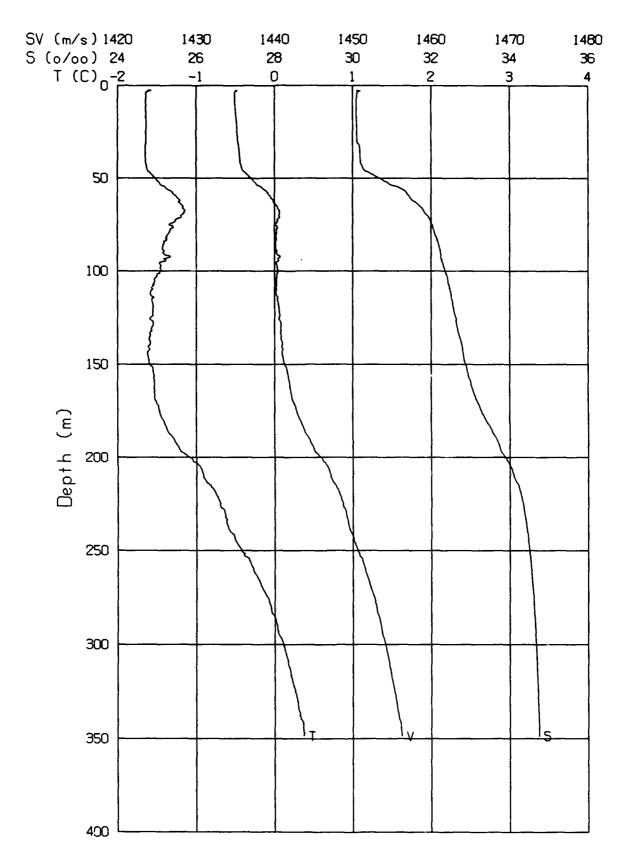


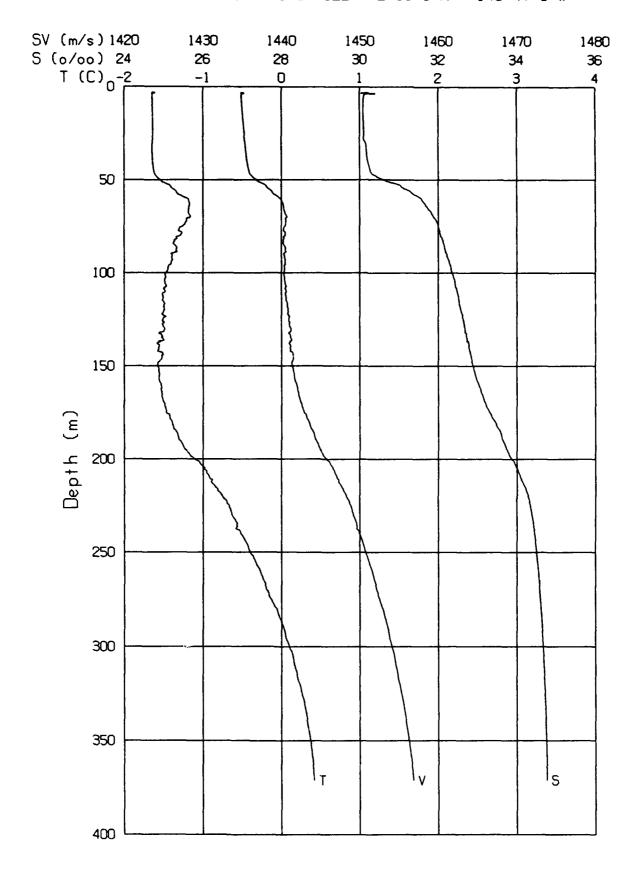


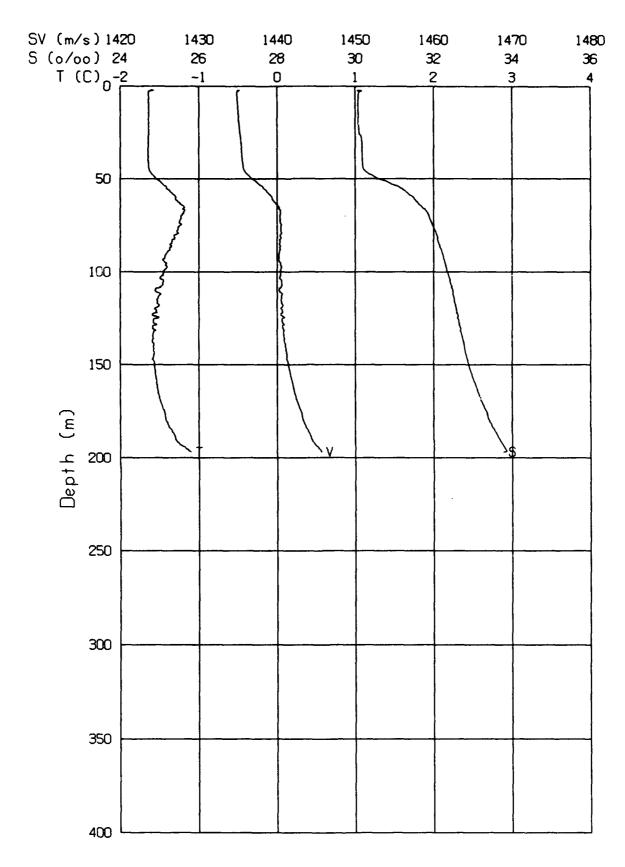


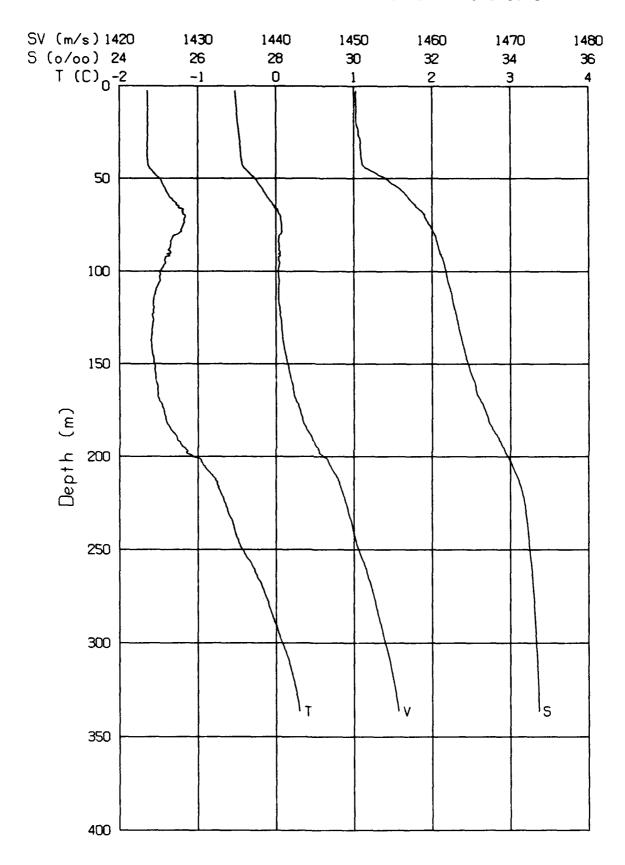


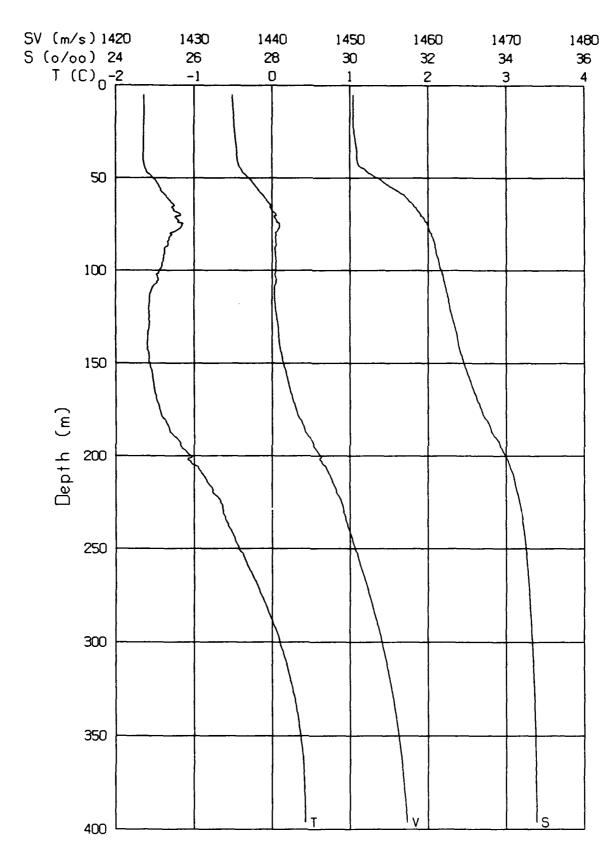


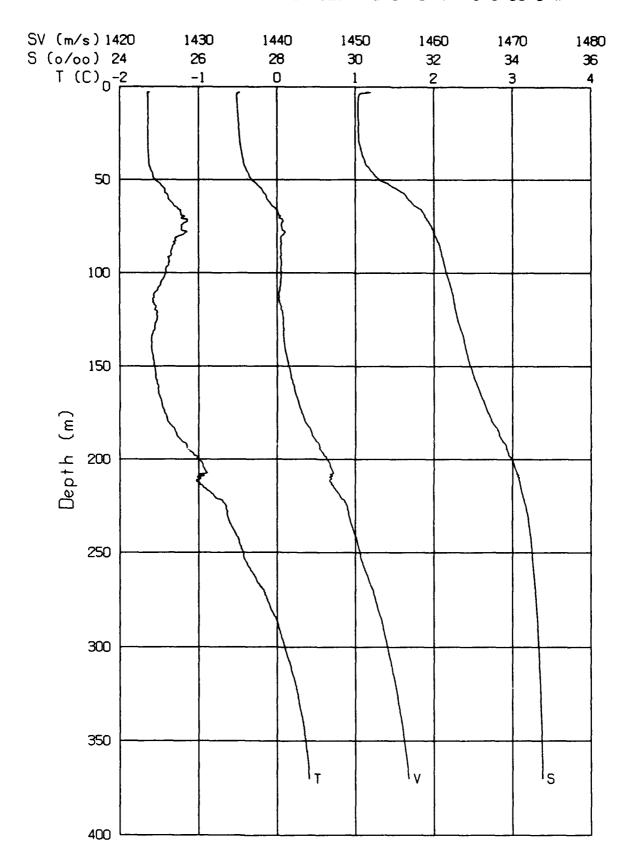


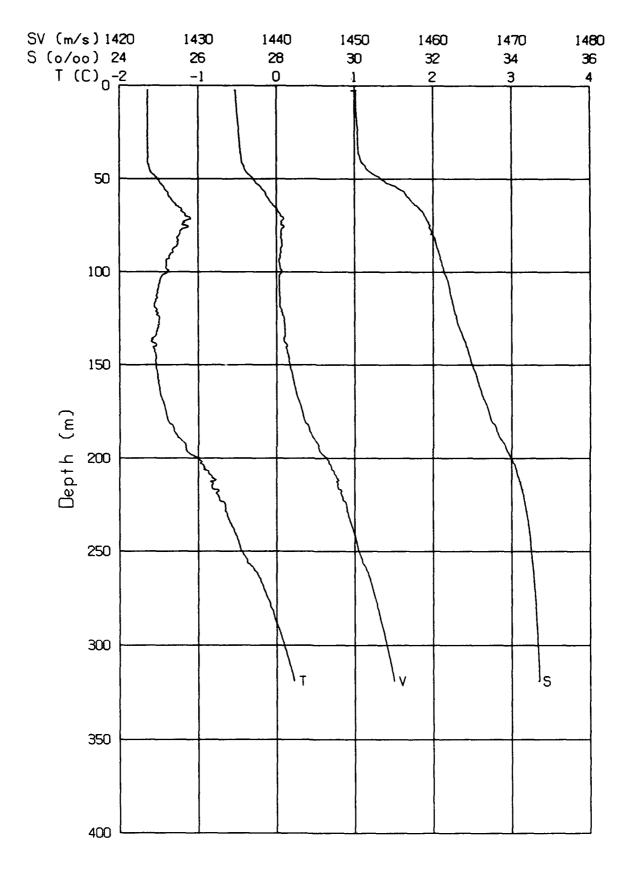


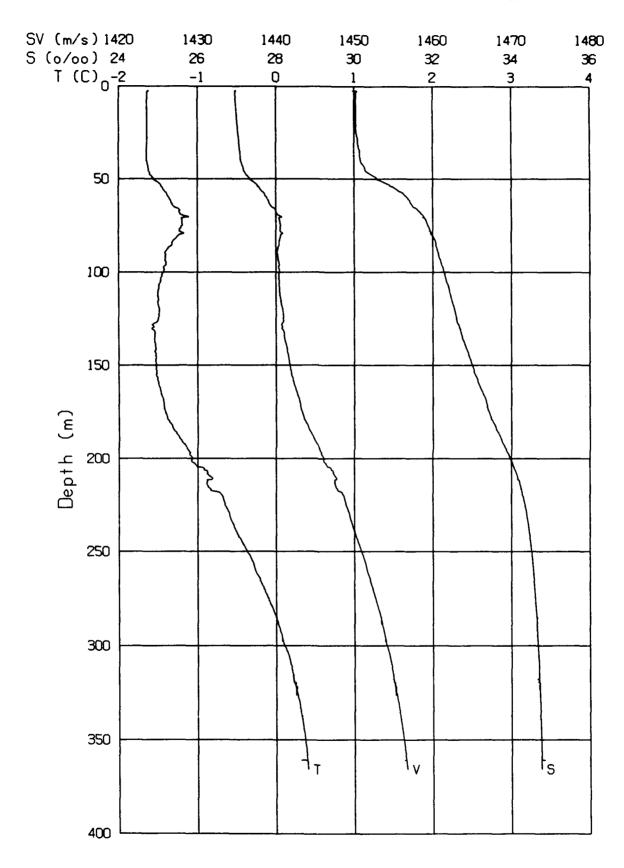


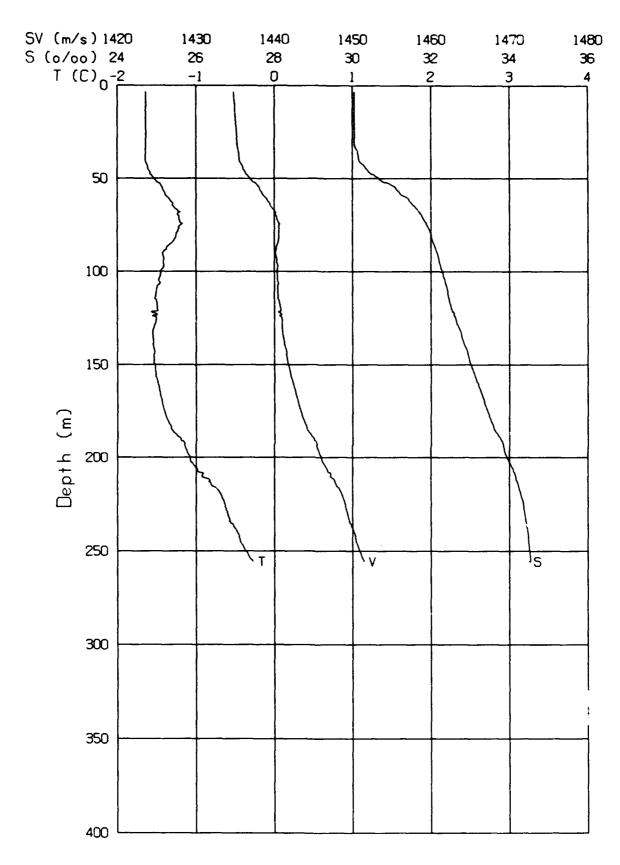


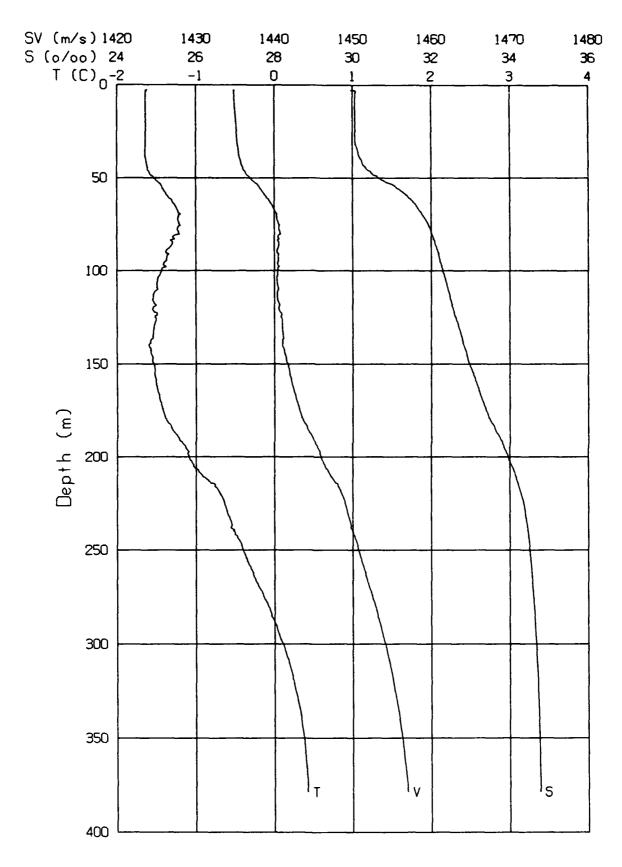


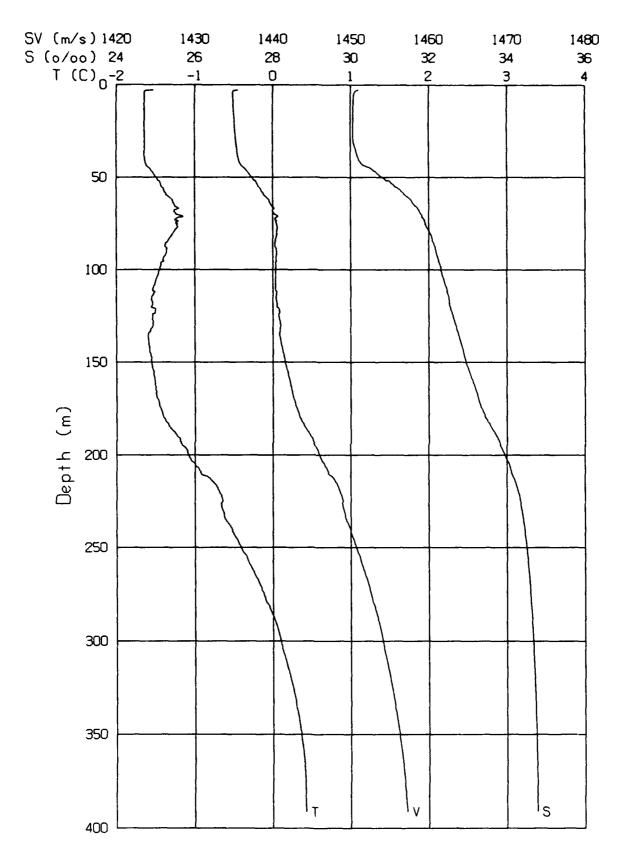


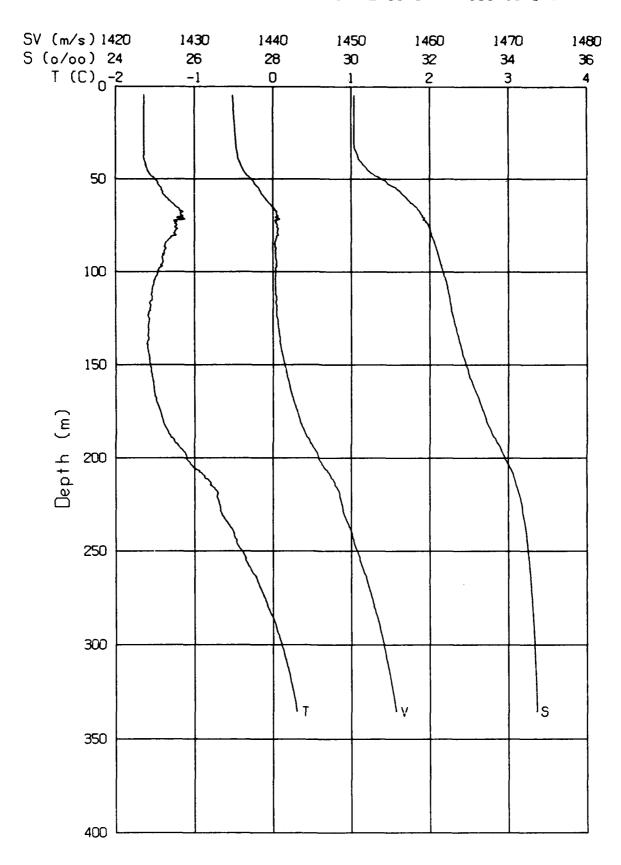


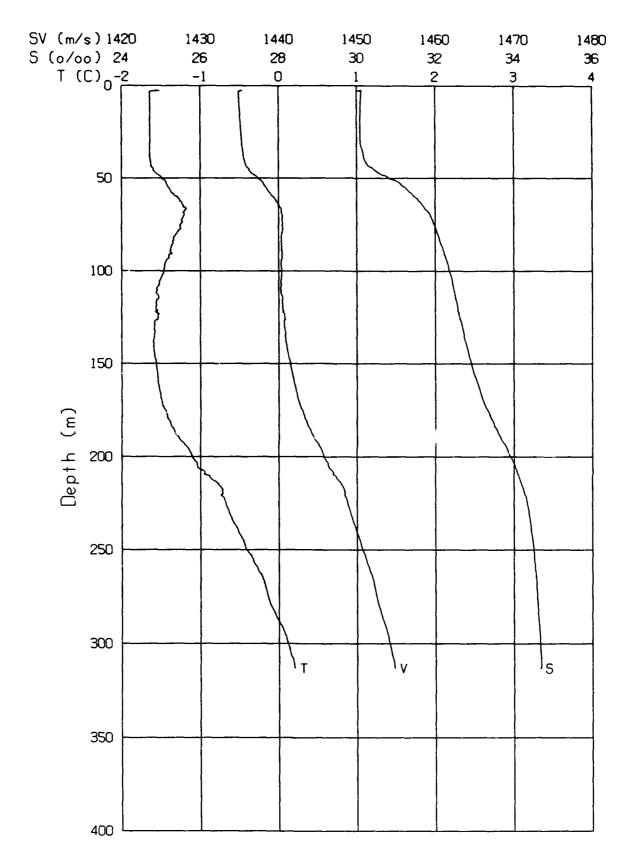


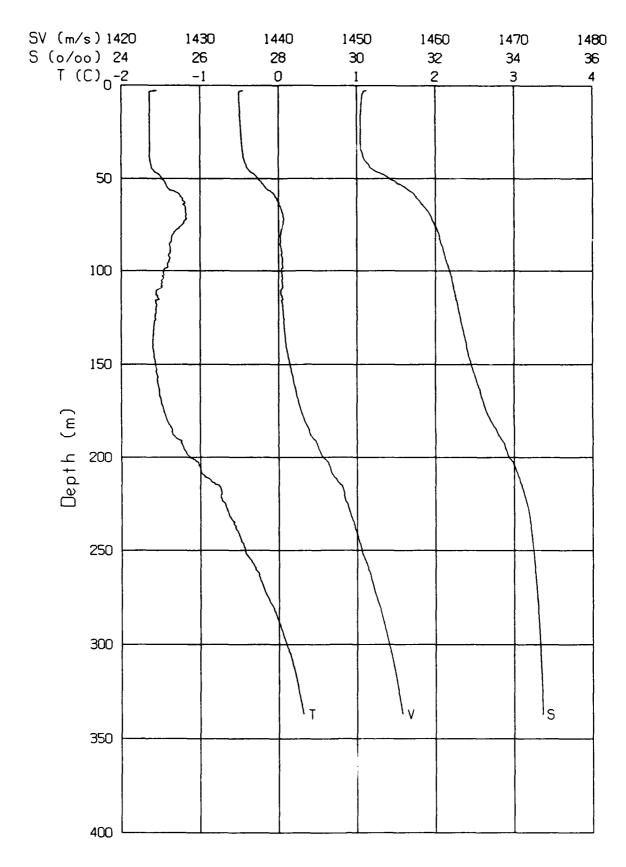


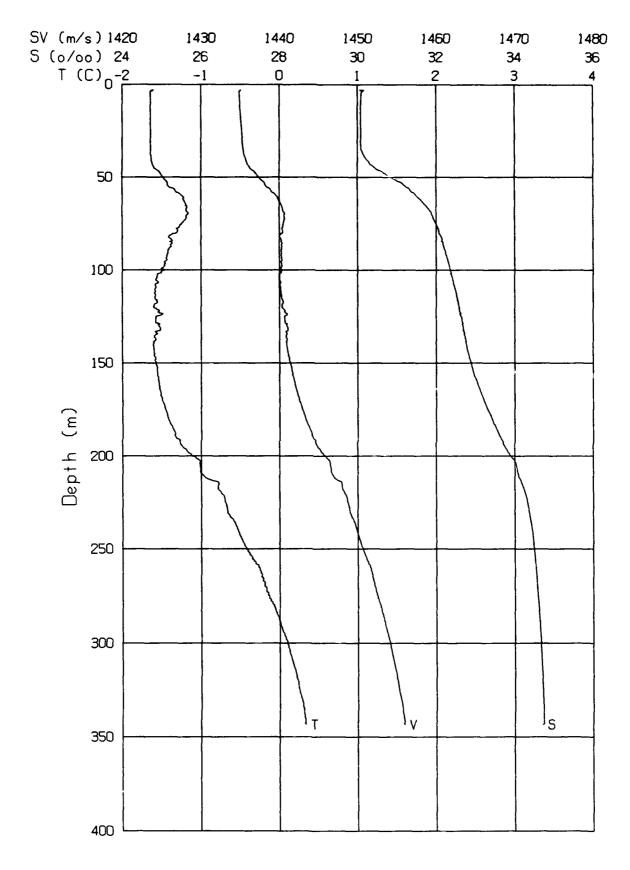


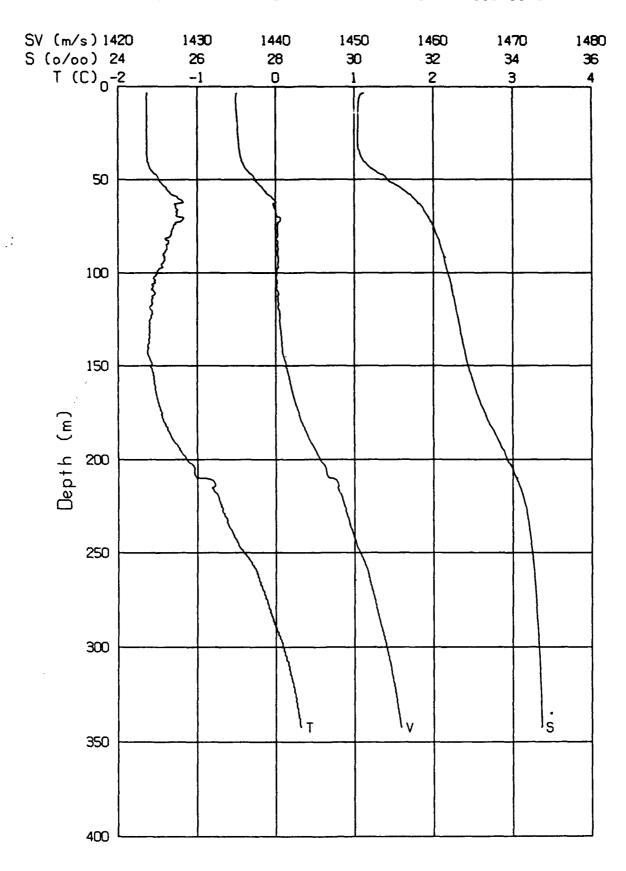


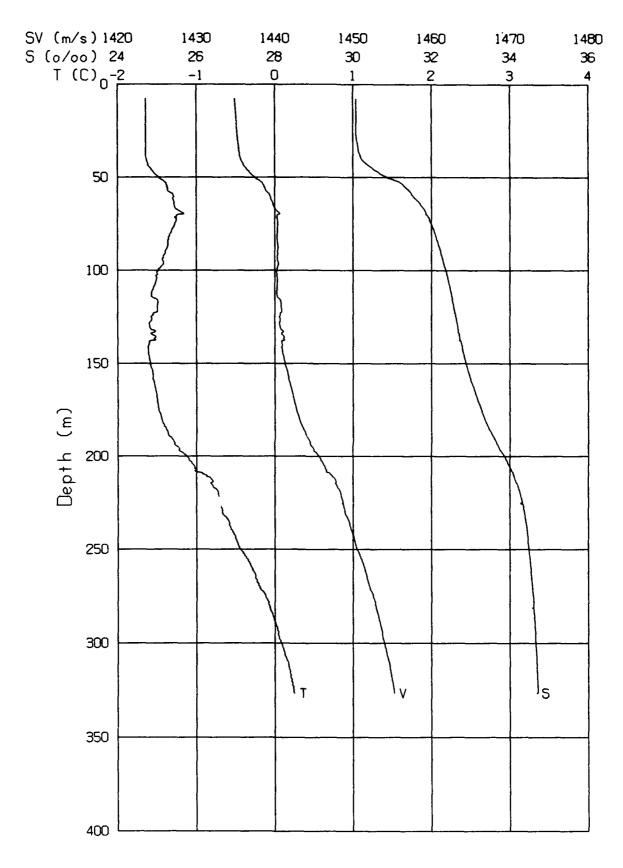


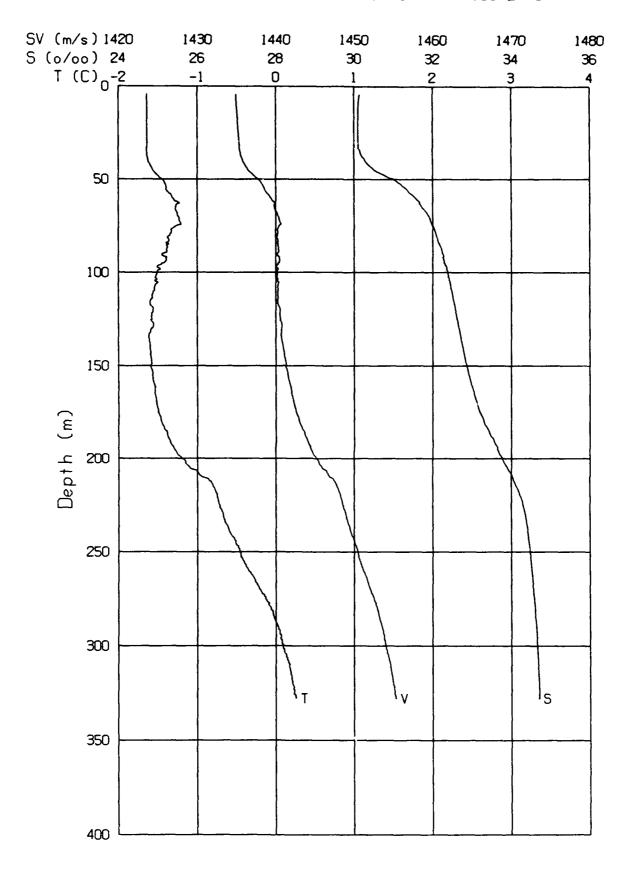


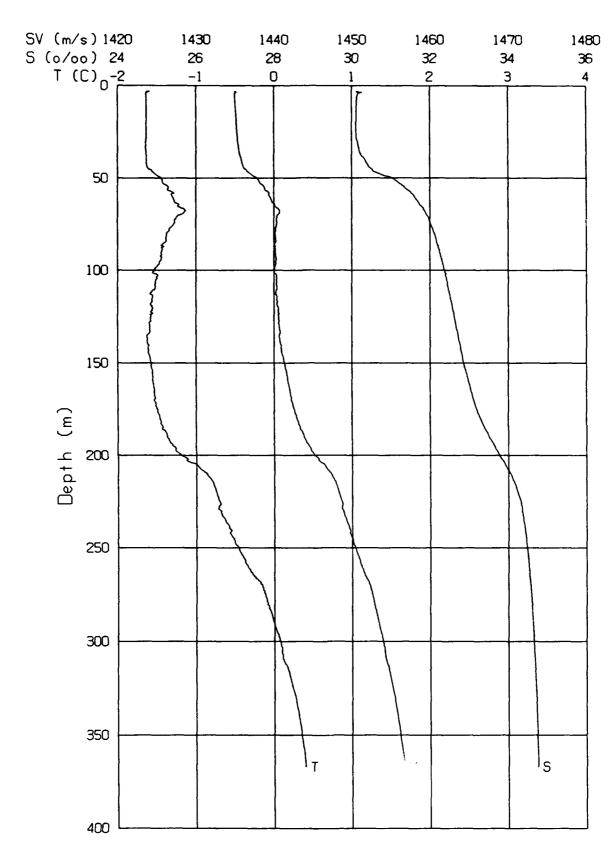


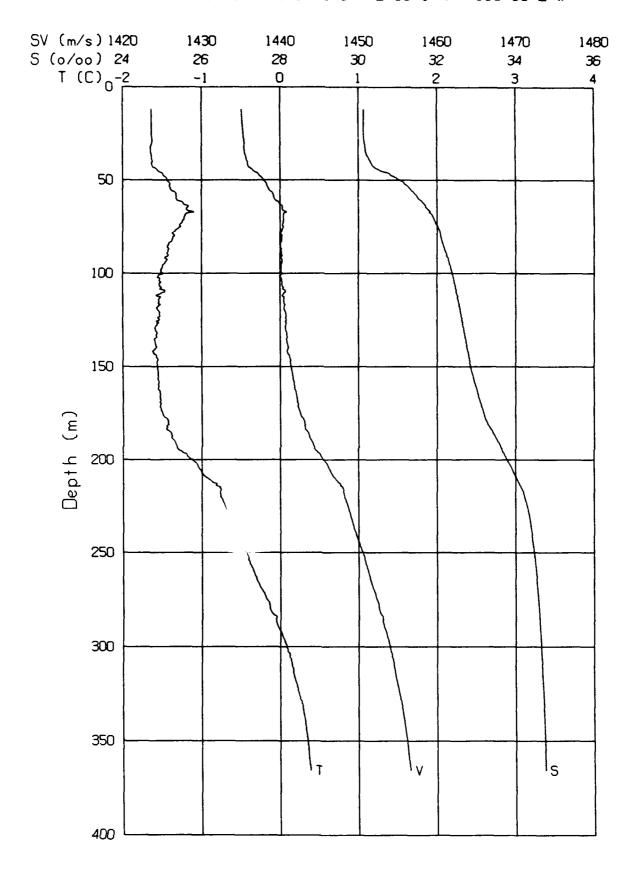




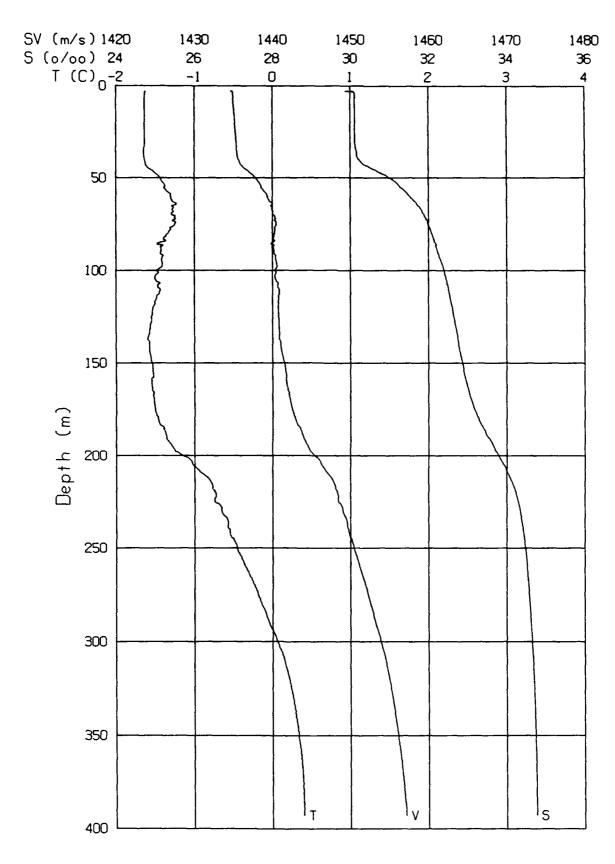


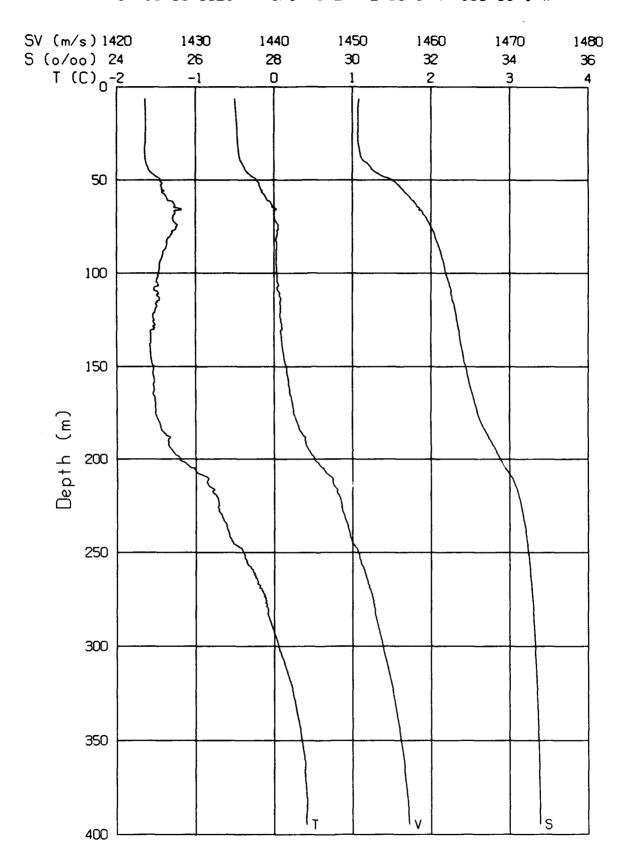


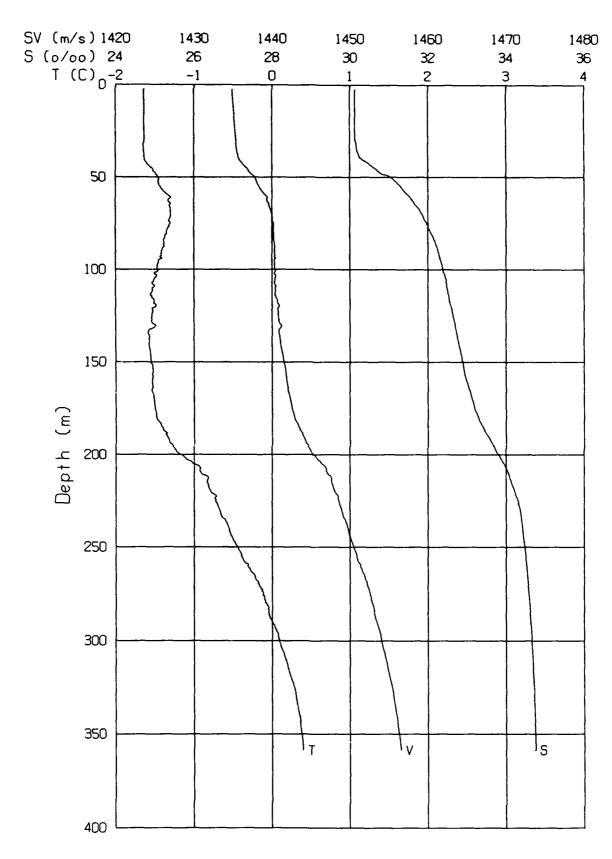




04-14-93 2134 CAST 041 72-38.3 N / 150-33.7 W

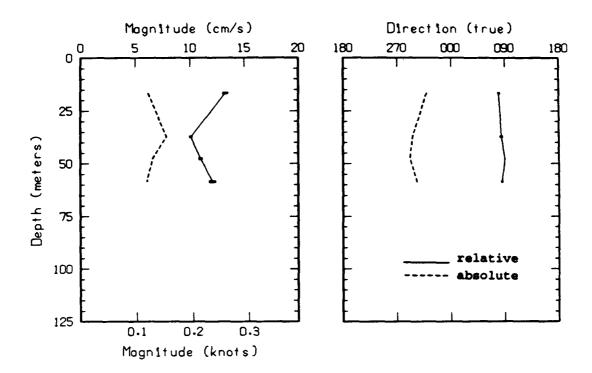




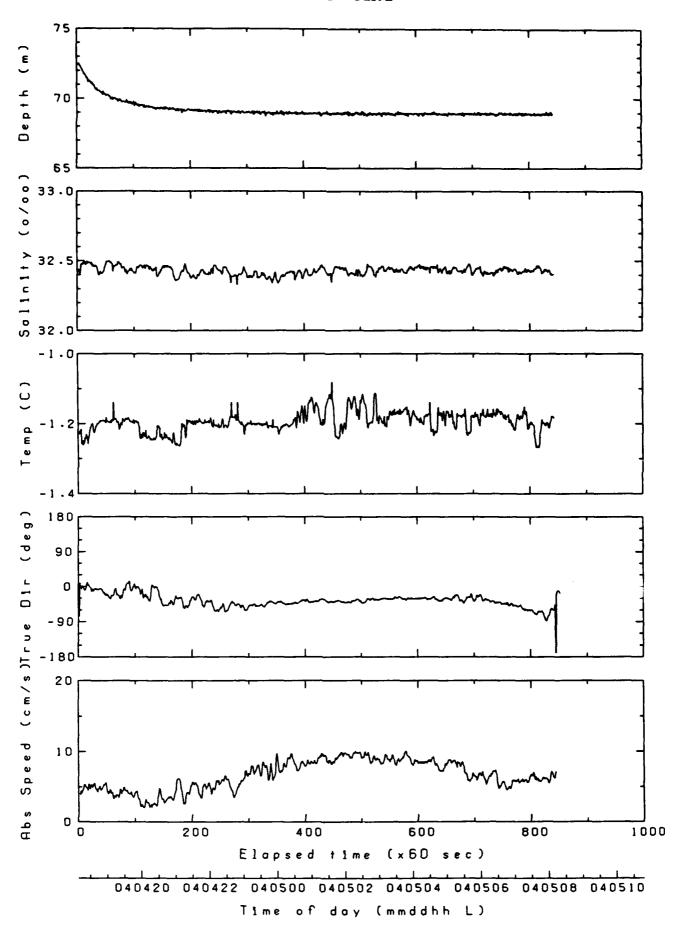


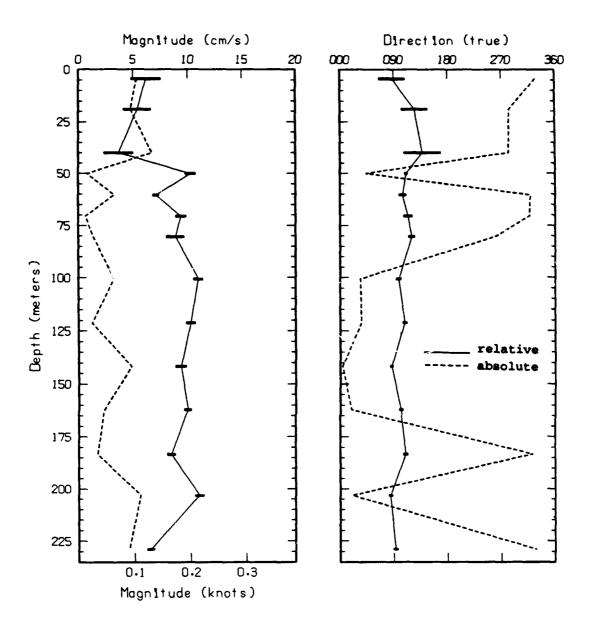
APPENDIX C
Current Meter Data

mmdd	hhmm (Local)	Cast	Floe Drift Ave (s)	cm/s	Dir	Comments
0401	1635	1	60	17.3	278	
0404	1810	2	60	12.0	281	Time series at 70 m
0405	1745	3	5	10.0	295	
0408	1145	4	5	4.8	254	
0409	1645	5	5	9.9	268	
0410	1815	6	5	4.9	280	
0411	1456	7	5	5.3	289	Time series at 225
0413	0915	8	5	5.0	286	Time series at 220



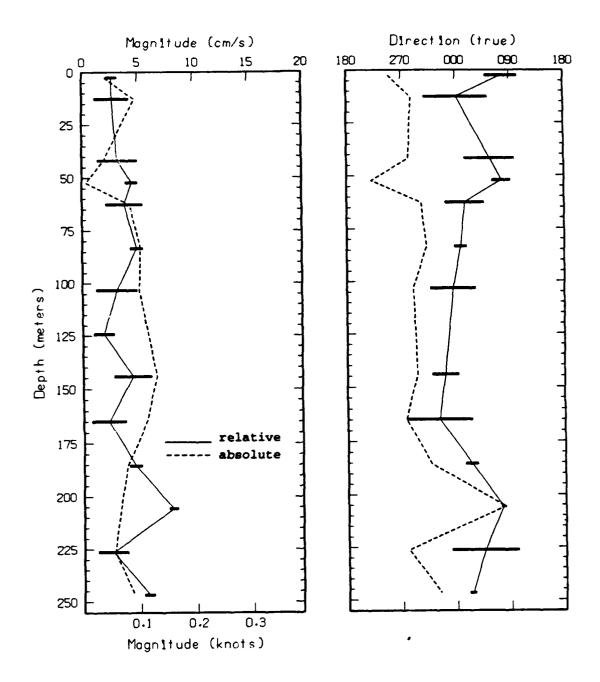
Magnetic bearing + 31 degrees = True bearing Floe drift speed = 17.3 cm/s Floe drift direction = 278 degrees True bearing





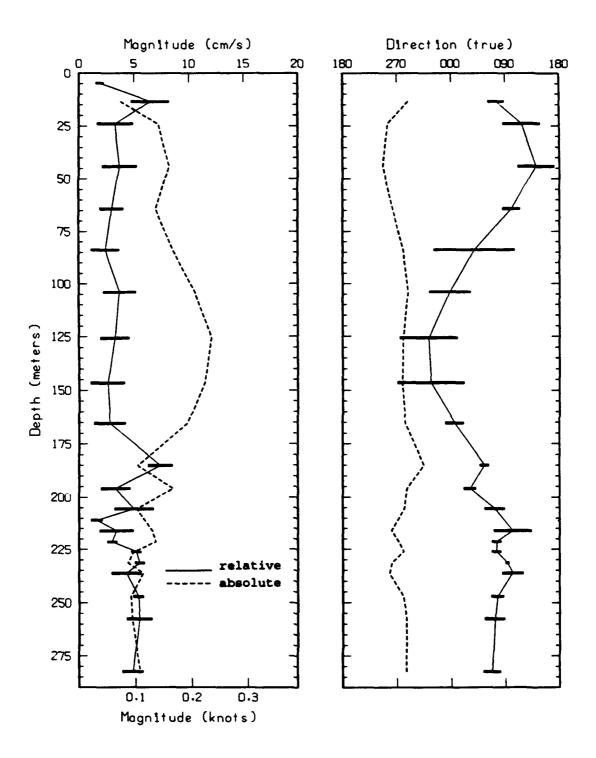
Magnetic bearing + 31 degrees = True bearing
Floe drift speed = 10.0 cm/s
Floe drift direction = 295 degrees True bearing

## 04/08 1145L Cast 4

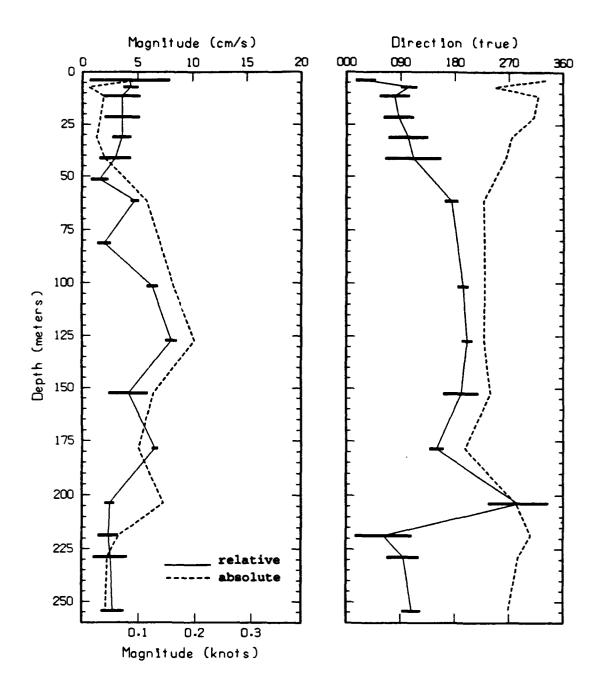


Magnetic bearing + 31 degrees = True bearing Floe drift speed = 4.8 cm/s Floe drift direction = 254 degrees True bearing

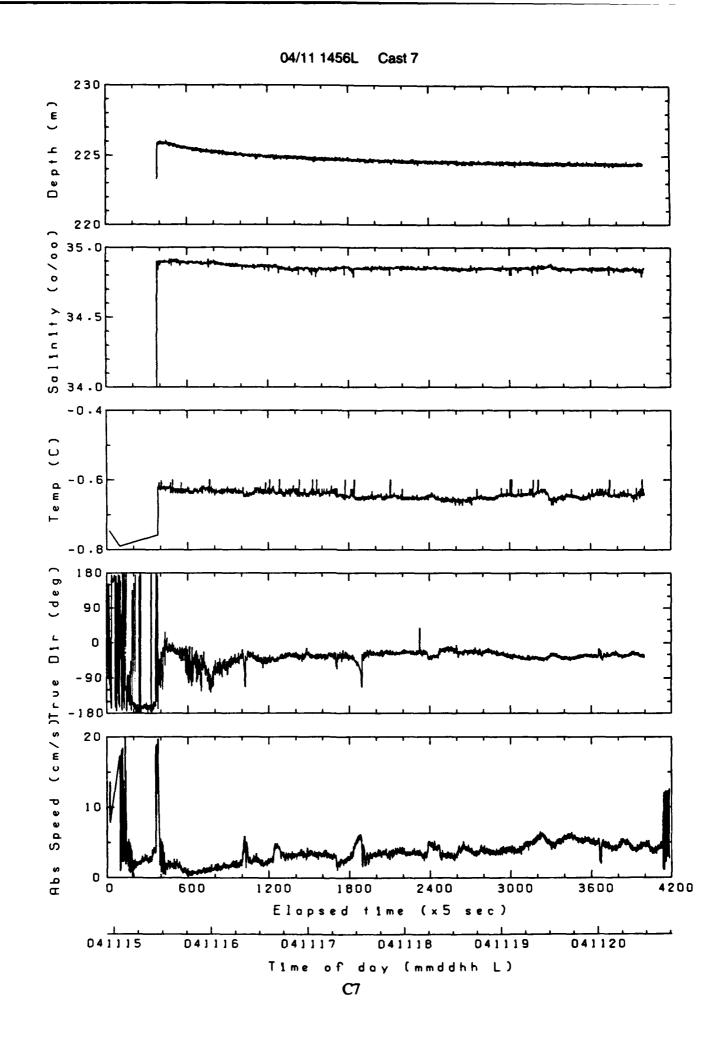
## 04/09 1645L Cast 5

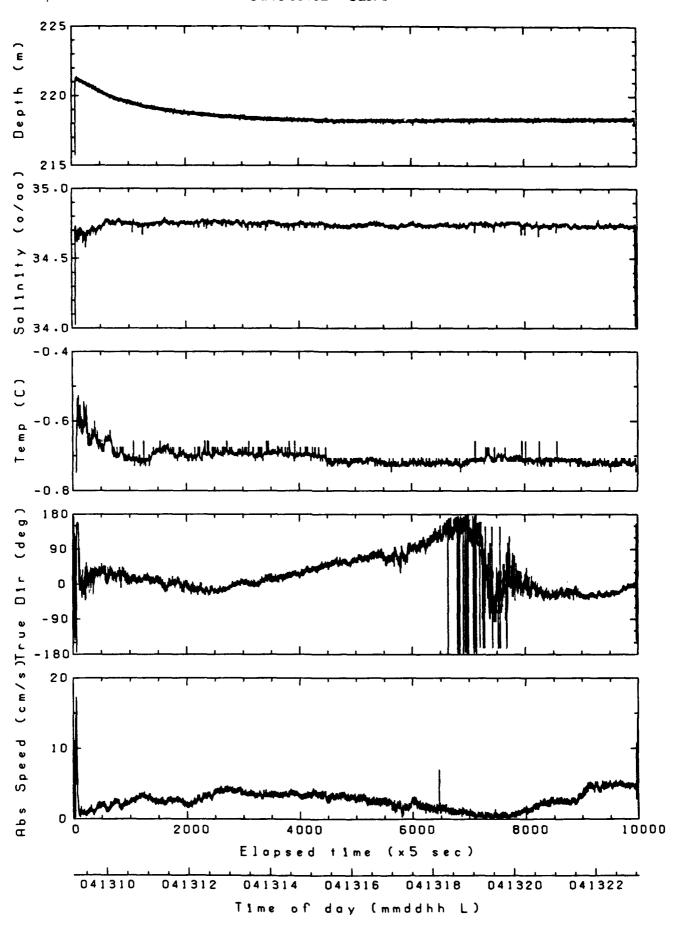


Magnetic bearing + 31 degrees = True bearing Floe drift speed = 9.9 cm/s Floe drift direction = 268 degrees True bearing



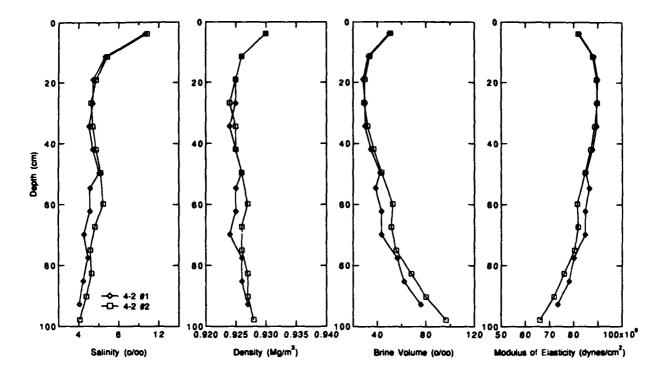
Magnetic bearing + 31 degrees = True bearing Floe drift speed = 4.9 cm/s Floe drift direction = 280 degrees True bearing



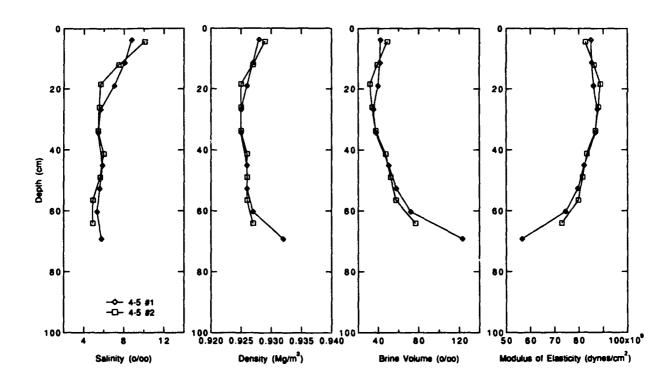


APPENDIX D

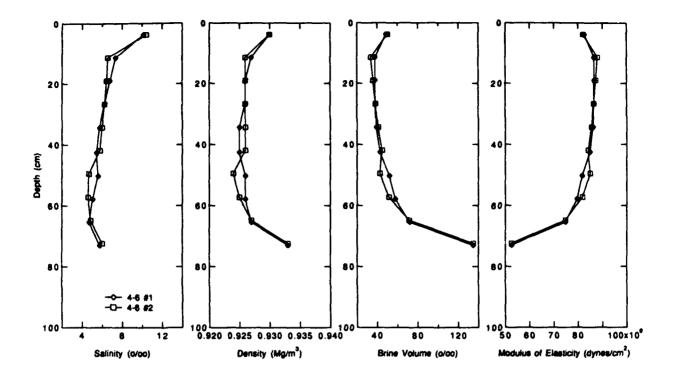
Core No.	Thickness (cm)	Surface Temp. (°C)	Comments
4-2 #1,2	102	-13.0	1.6 m apart
4-5 #1,2	74	-12.7	> 2 m apart
4-6 #1,2	76	-12.7	> 5 m apart
4-11 #1	84	<b>~13.1</b>	•
#2,3	<b>8</b> 5	-13.6	0.5 m apart, >10 m from #1
4-12 #1,2		****	Refrozen melt ponds



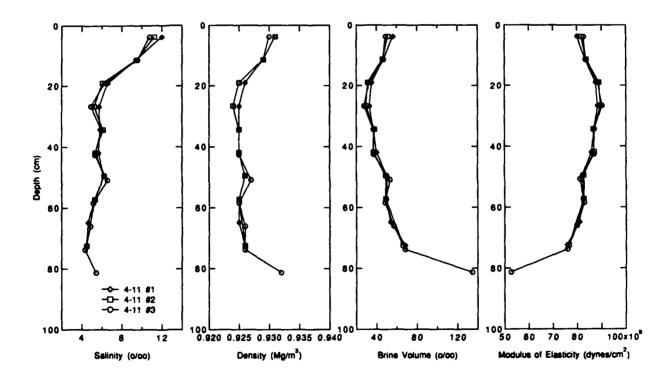
4/2 - #1, #2



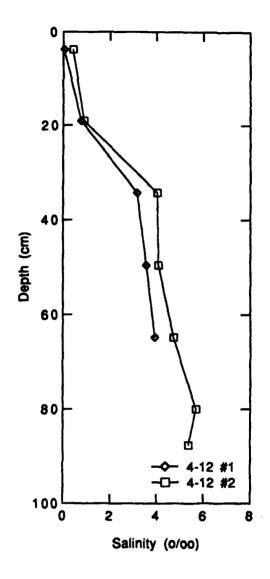
4/5 - #1, #2



4/6 - #1, #2



4/11 - #1, #2, #3



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F. Symons, Jr. J. Kisenwether

APL-UW

K. Aagaard A. Brookes W. Felton F. Karig J. Luby R. Stein T. Wen K. Williams

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	Data analysis is limited because t								
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